

Topic	Examples of interactive data analysis of seismic records using the SEISAN software
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Version	April 2013; DOI: 10.2312/GFZ.NMSOP-2_IS_11.6

Note: Figure numbers followed in this text by ?? relate to the NMSOP editions 2002 (printed) and 2009 (website). They will be changed when the revised Chapter 11 of NMSOP-2 becomes available.

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1 Introduction

The purpose of this information sheet is to illustrate the steps involved when analyzing seismic records by means of the SEISAN software. Examples are given for local, regional and teleseismic earthquakes. This is done by giving representative screen plots together with explanatory texts that follow the figures.

For details on the SEISAN software the reader is referred to section 11.4.2 (Data Analysis and Seismogram Interpretation - Software for routine analysis – SEISAN) and the references therein. Details of the software commands are not given here, and the reader is referred to the SEISAN manual for details, which can be downloaded either from NMSOP-2 via the link ‘Download programs and files’ or from <ftp://ftp.geo.uib.no/pub/seismo/SOFTWARE/SEISAN/>.

The events are included with the sample data that is available with SEISAN (from version 9.2).

The main background material provided by NMSOP-2 (<http://nmsop.gfz-potsdam.de>), explaining the methodology behind this information sheet, is:

- CHAPTER 2: Seismic Wave Propagation and Earth Models
- CHAPTER 3: Seismic Sources and Source Parameters
- CHAPTER 5: Seismic Sensors and their Calibration
- CHAPTER 6: Seismic Recording Systems
- CHAPTER 11: Data Analysis and Seismogram Interpretation
- IS 2.1: Standard nomenclature of seismic phases
- IS 3.3: The new IASPEI standards for determining magnitudes from digital data and their relation to classical magnitudes
- IS 11.1: Seismic source location
- IS 11.4: Tutorial for consistent phase picking at local to regional distances
- DS 2.1-2.4: Record examples for local, regional and teleseismic earthquakes and explosions

Note that there are in the following some, only apparent, differences between SEISAN internal nomenclature and procedures and those recommended by the IASPEI standards for measuring magnitudes which do not contradict. While SEISAN mb and IASPEI mb are identical, SEISAN mB corresponds to IASPEI mB_BB. Similarly holds that SEISAN Ms = Ms_20 and MS = Ms_BB. Moreover, IASPEI standards recommend to measure mB_BB and Ms_BB on unfiltered broadband velocity traces, however accompanied by the statement that their passband should cover at least the period range within which these two magnitudes should be measured. SEISAN, however, always deconvolves any given actual record so as to simulate for mB_BB a velocity-proportional broadband response in the period range between 0.2 and 30 s and for Ms_BB a response in the period range between 3 and 60 s. This allows in SEISAN to read the broadband magnitudes also on short-period records, provided that the SNR is sufficient to allow a broadband velocity restitution.

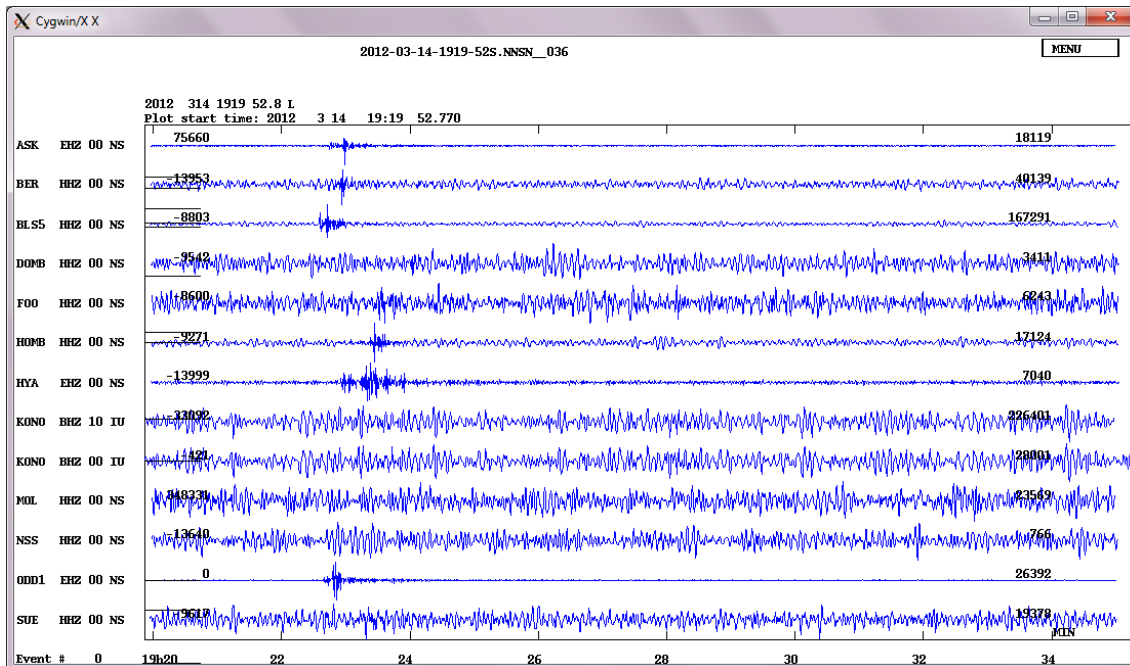
2 Analysis of a local earthquake in southern Norway recorded between 50 km and 650 km epicentral distance

The earthquake chosen here as example occurred in southern Norway and was recorded on the Norwegian National Seismic Network operated by the University of Bergen. The event was recorded on a mix of short-period and broadband seismic stations. It was automatically detected and then interactively processed.

The earthquake parameters are:

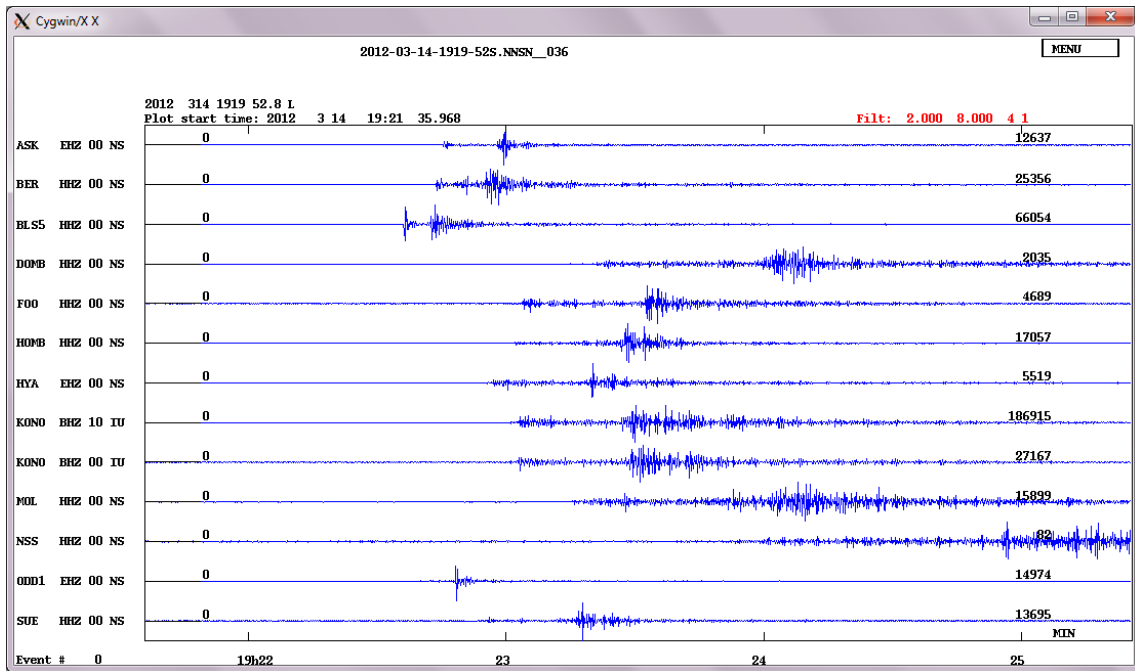
Date:	14 Mar 2012
Origin time:	19:22:27.9 (UTC)
Latitude:	59.526°N
Longitude:	5.615°E
Depth:	13.9 km
ML:	3.4

We now show the step-by-step procedure to analyze the earthquake.

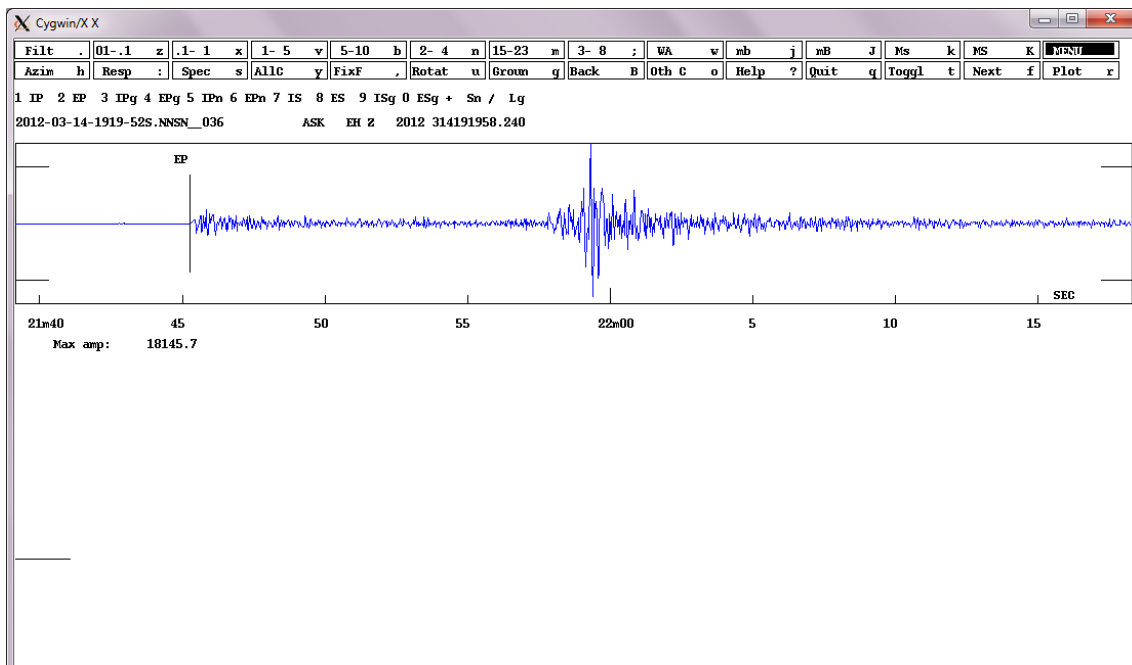


Vertical-component traces, unfiltered: Our starting point is a waveform file that is produced by an automatic detection system. No additional information is given, although, depending on the automated processing system, additional information including type of event, location and magnitude may already be available. It is our first task to identify the type of event that is recorded. If the detection is not false, the initial choices are between local, regional and teleseismic event. It may also be possible to identify if the event is likely an explosion based on the characteristics of the waveforms and the pattern of phase arrival times. The above multi-trace plot shows records of a small local earthquake, characterized by the short signal duration of less than about 1 min and a relatively high frequency signal content.

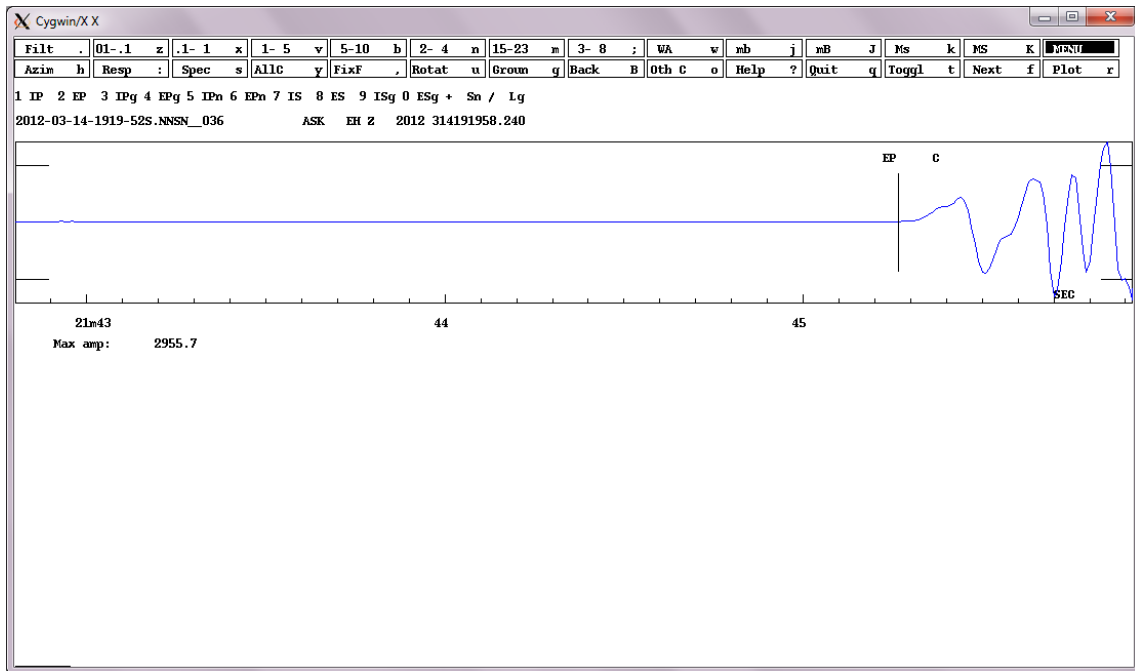
The event is recorded on 13 seismic stations. The numbers on the SEISAN multi-trace plot gives the DC of individual traces on the left and the maximum zero to peak amplitude on the right. The units are counts unless instrument correction is used to compute displacement, velocity or acceleration with the units nm (nanometers = 10^{-9} m), nm/s and nm/s², respectively. Station and component codes are given left of each trace together with the location code and network code. For the first trace, the station is ASK, the component is EHZ (short-period seismometer, with high sampling rate of more than 80 samples per second), the location code is 00 (used to indicate differences between equipment that gives the same component code at the same site, for example an STS2 at 40sps is labeled BHZ.00 and an STS1 at 20sps is labeled BHZ.10) and the network is NS (indicating who operates the station, in this case NS = Norwegian National Seismic Network). The event is well visible on the short-period seismometer records, which have the component code EHZ. However, it is not seen on the broadband traces, which are labeled HHZ, as this is a small earthquake where the signal content is mostly above 0.5 Hz.



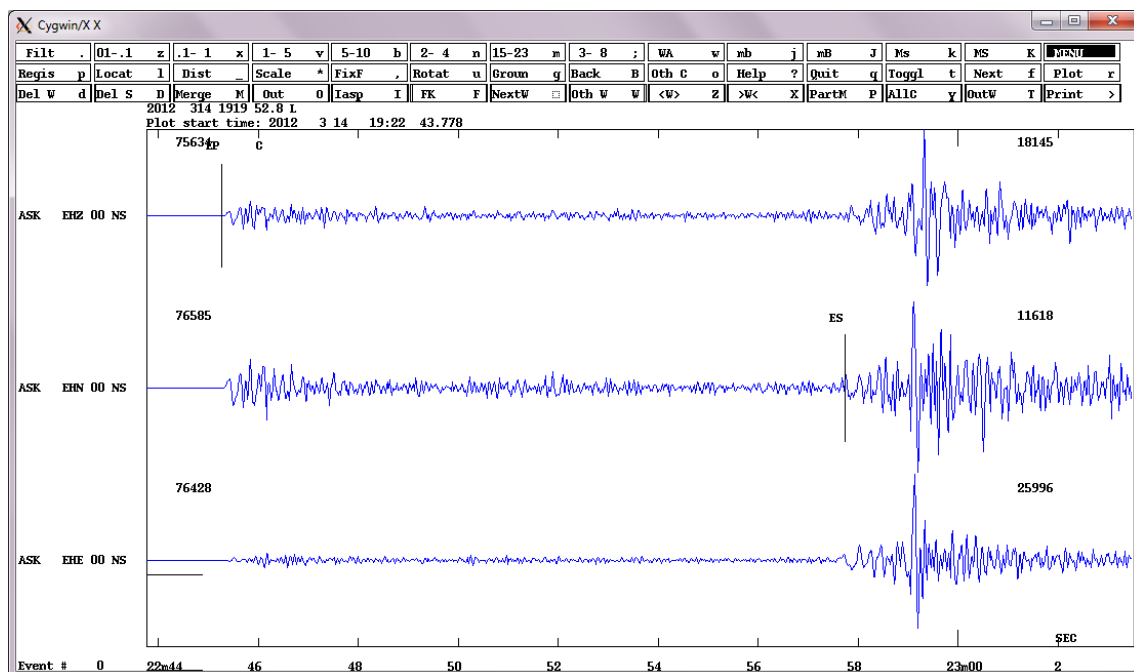
Vertical-component traces, filtered 2-8 Hz: Applying this high-frequency band-pass filter, which suppresses the more long-period microseismic noise in the broadband records, the earthquake is now clearly seen on all the traces. As compared to the previous plot, the time window of signal arrivals has been cut out and the time scale been stretched. With distinct primary P- and secondary S-wave arrivals in this frequency band it is very clear that this is a local earthquake. We can see from the arrival times that the earthquake was closest to station BLS5. In SEISAN we then put this event into the database and start the processing.



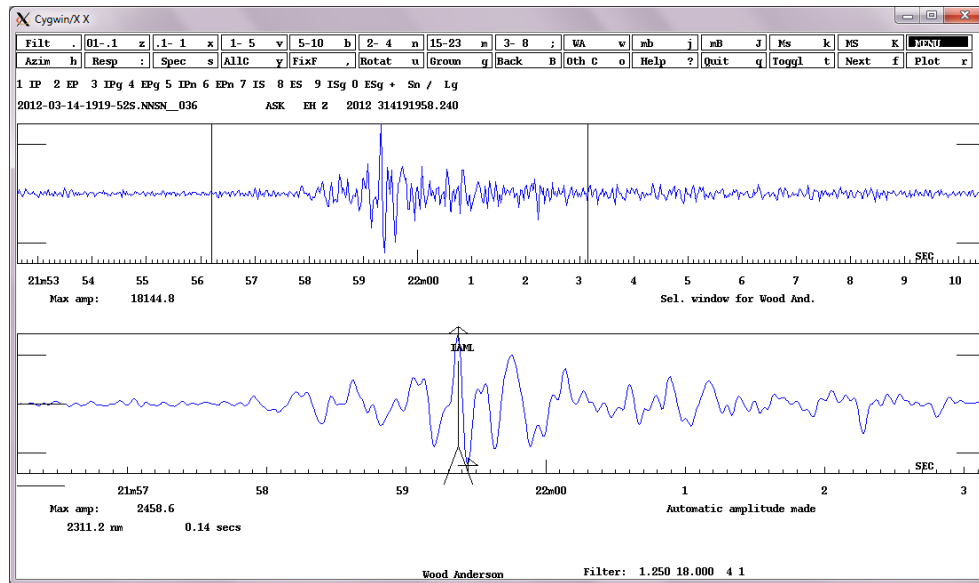
Single trace plot of station ASK at epicentral distance of 109 km, unfiltered: In SEISAN, we mostly do the phase picks on single trace plots, like shown here, or on three-component trace plots. For this recording on a short-period seismometer it is not necessary to filter the data in order to pick phase arrivals. The P arrival is marked.



Single-trace plot of station ASK at epicentral distance of 109 km, unfiltered: This is the same station as on the previous plot, but this time we have zoomed in on the P arrival. It is now possible to read the P arrival time more accurately, and we can also identify the polarity of the first arrival, which is compression (upward motion on the vertical component) in this case. The phase arrivals will be used for the earthquake location (see IS 11.1) and the polarity can be used to determine the fault plane solution (see also EX 3.2).

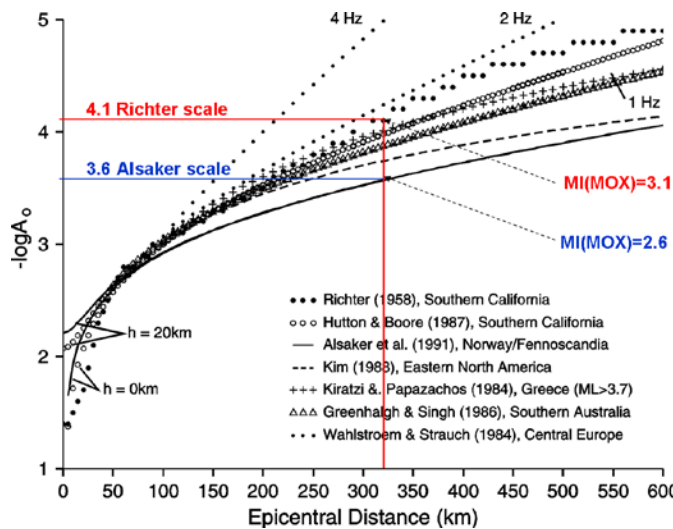


Three-component trace plot of station ASK at epicentral distance of 109 km, unfiltered: We normally read the S arrival on one of the two horizontal channels and not on the vertical as SP conversions may arrive prior to S and result in an incorrect reading of the onset time.

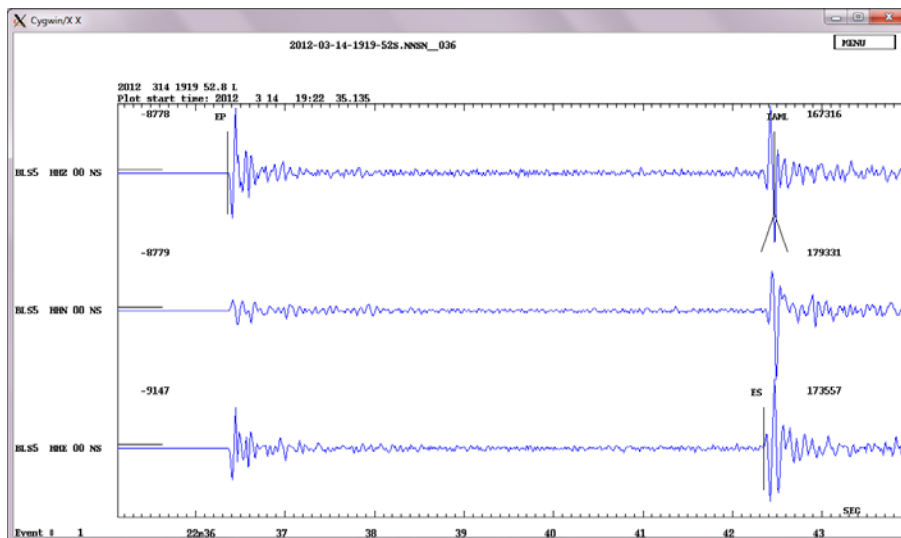


Single-trace plots of station ASK at epicentral distance of 109 km, Wood-Anderson filter: To read the ML amplitude, the instrument response is removed and a filter equivalent to the revised Wood-Anderson seismograph response according to Uhrhammer and Collins (1990) (see also IS 3.3) is applied. The bottom trace is a zoomed-in version of the area between the two vertical lines on the top trace. In SEISAN the amplitude is given in units of nanometer. It is measured peak-to-adjacent trough of the maximum deflection of the seismogram trace, where peak and trough are separated by one crossing of the zero-line (sometimes described as “peak-to-peak amplitude”) but stored in the database is peak-to-peak/2 according to the IASPEI (2013) standard. This maximum amplitude is normally read from the S or Lg waves. It is stored as the IAML amplitude phase. ‘I’ stands for IASPEI standard, assuring that the standard Wood-Anderson filter response and the recommended amplitude measurement procedure has been applied (see IS 3.3), and ‘A’ stands for displacement amplitude in nm. IASPEI standard for ML is to measure A on horizontal components. In Norway, however, it is routine practice to use the vertical component for the amplitude reading of ML and the scale given by Alsaker et al. (1991; see also DS 3.1, Tab. 2). This scale was derived for Norway and scaled so that its results match at 60 km distance with those of the horizontal component California standard formula for equal IAML input (see next figure and recommendations in IS 3.3). The default ML formula implemented in SEISAN is that of Hutton and Boore (1987; see next figure and formula in DS 3.1, Tab. 2), on which also the IASPEI standard formula for ML is based (see IS 3.3). If, however, this default procedure is used in other regions with other attenuation conditions then the calculated MI values may differ from standard ML. DS 3.1. Table 2, offers several MI calibration functions for other regions together with comments on their scaling, components to be used and the distance range of their applicability. You may implement in SEISAN either your own properly to the ML standard scaled local calibration function, or, if not yet available, one of those which you believe to match best with the seismotectonic conditions in the area. Note, however, that local MI data based on not yet tested and proven standard scaling should not be published with the nomenclature for standard ML. That is, why we use in NMSOP MI as the general nomenclature for local magnitude, and ML only for MI that has been properly scaled to the new standard. Otherwise, users of local magnitude data do not know whether or not they have been correctly scaled according to the now globally set standard. E.g., according to Braunmiller et al. (2005) (see also section 3.2.9.6 in the forthcoming new Chapter 3) the MI scales used by various agencies in Southern Germany, France, Italy and Switzerland yield results for equal input data that differ on average between 0.2 to 0.6 m.u. when scaled to Mw.

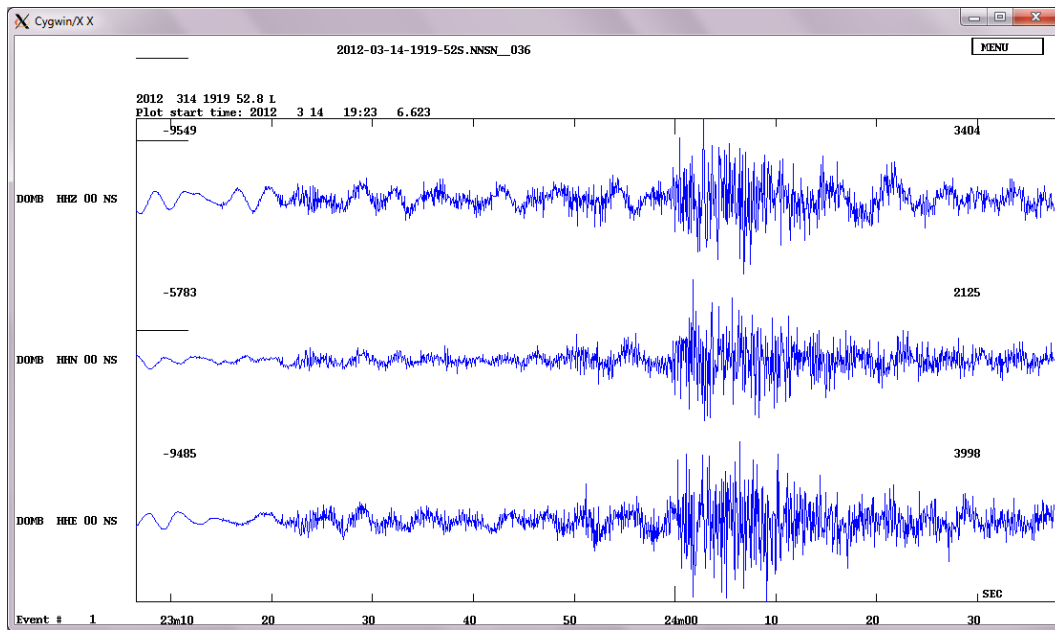
Some examples for other regional MI calibration functions have been plotted in the diagram below. Note that the Alsaker et al. (1991) scale (lowermost full line) for the continental shield areas of Scandinavia and Eurasia reveals significantly lower attenuation when compared with Southern California calibration curves of both Richter (1935) and Hutton and Boore (1987) (see full dot and open dot curves in the Figure below). Accordingly, since properly scaled to the latter at about 60 km epicentral distance, the Alsaker et al. calibration curve yields at several 100 km distance for equal amplitude input data already ML values that are several tenths of magnitude units smaller than those derived by applying the original Richter scale. This difference increases to 1.7 m.u. at a distance of 1500 km, up to which the Alsaker et al. (1991) scale is defined (although strictly Hutton and Boore (1987) is limited to 600 km).



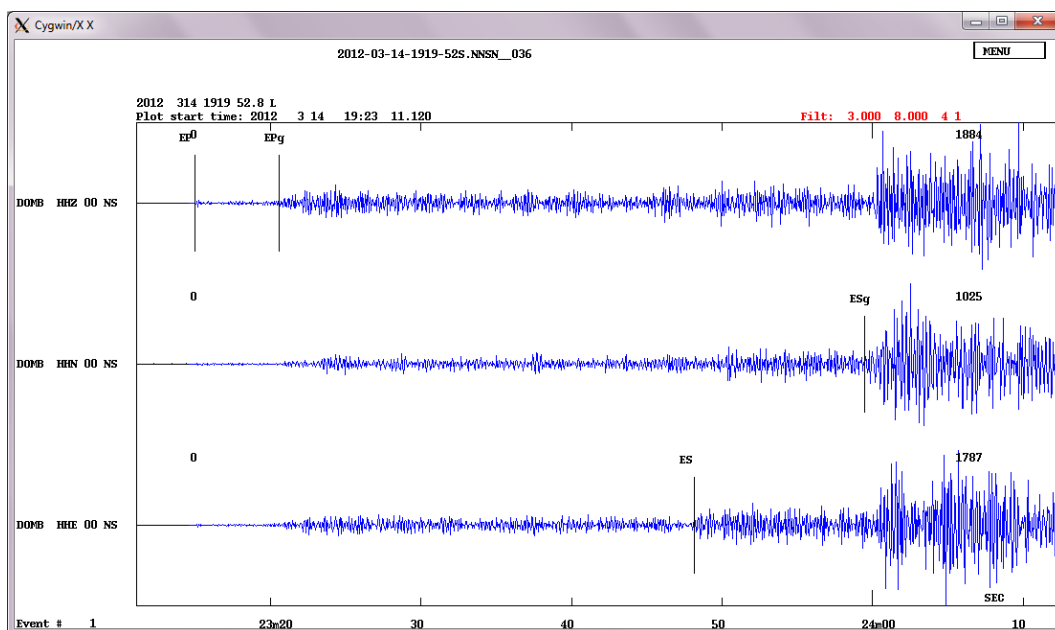
We now repeat the same procedure of measuring IAML for the other stations, however show only examples from some selected stations to explain specific aspects of the data processing.



Three-component traces of station BLS5 at epicentral distance of 49 km, unfiltered: Plotted are the three-component broadband traces from the nearest station. P and S arrivals, and the ML amplitude were read. In this case the SNR was high enough so that it was not necessary to filter the broadband recordings to identify the phase arrivals.

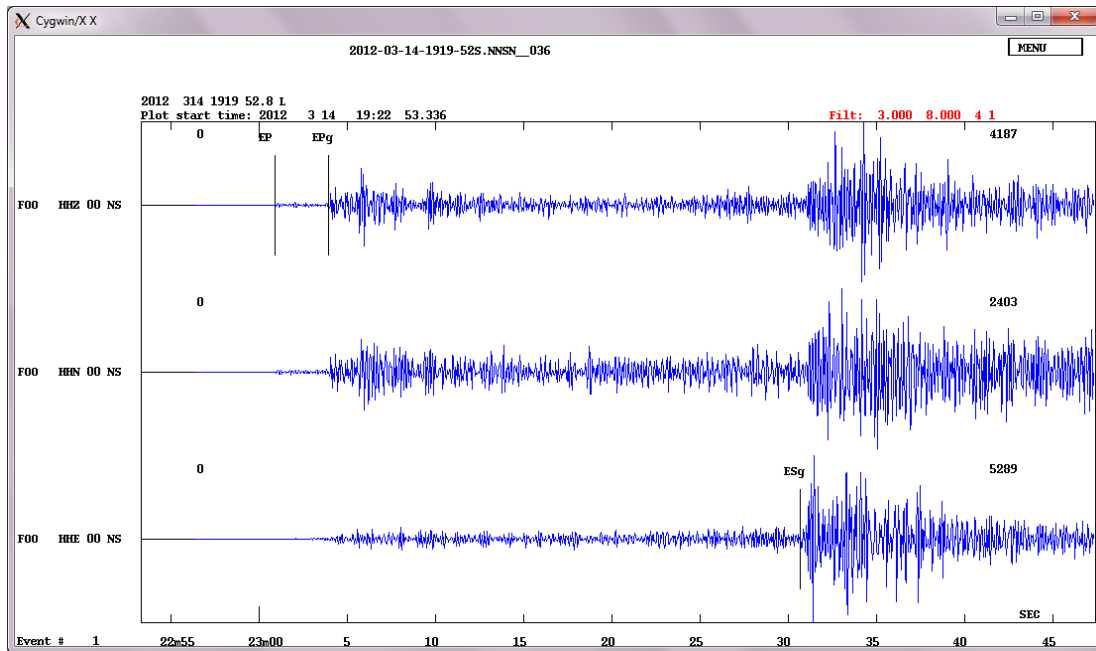


Three-component traces of station *DOMB* at epicentral distance of 342 km, unfiltered: Going to larger distances, the signal amplitudes at higher frequencies decrease relative to the more low-frequency microseismic noise, which peaks around 5 seconds. Here, we would normally filter the traces to read the phase arrivals.

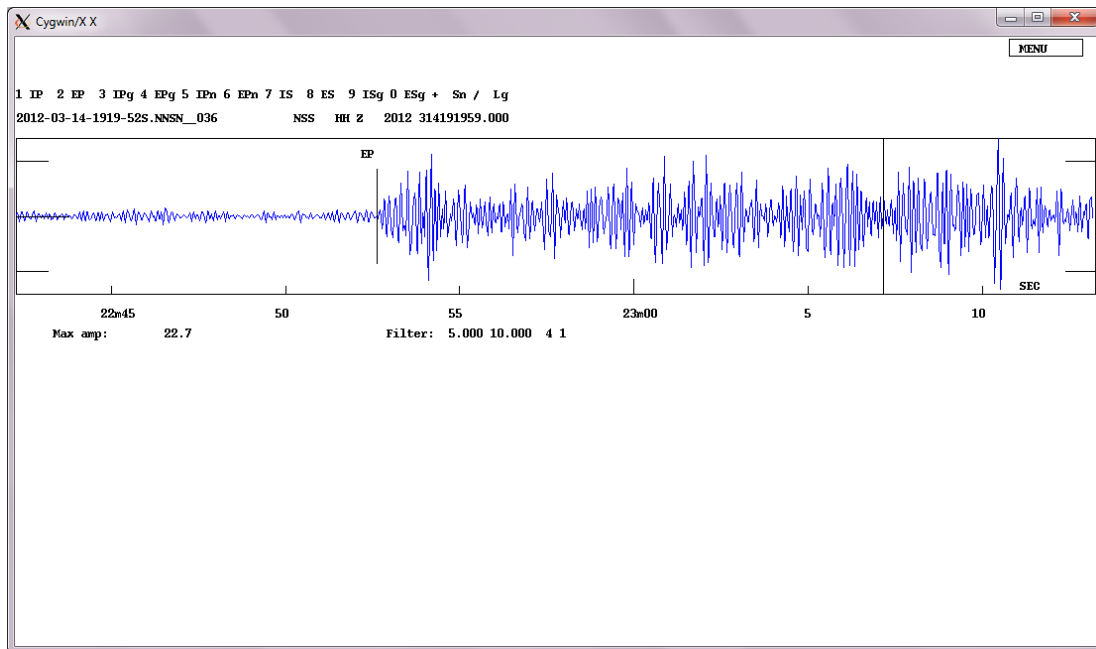


Three-component traces of station *DOMB* at epicentral distance of 342 km, filtered 3-8 Hz: For this station, it is possible to read besides the larger amplitude P and S arrivals, termed Pg and Sg, additionally the commonly weaker but earlier arrivals Pn and Sn. In SEISAN we generally mark the first P- and S-wave arrivals as P and S, respectively. The location program then determines which phase arrives first. Sometimes, Sn (here marked as ES) is weak and not recognizable in the signal-generated noise of scattered P waves. Then one has to mark Sg as such, since the location program otherwise would use it as the first arriving S (Sn in this case).

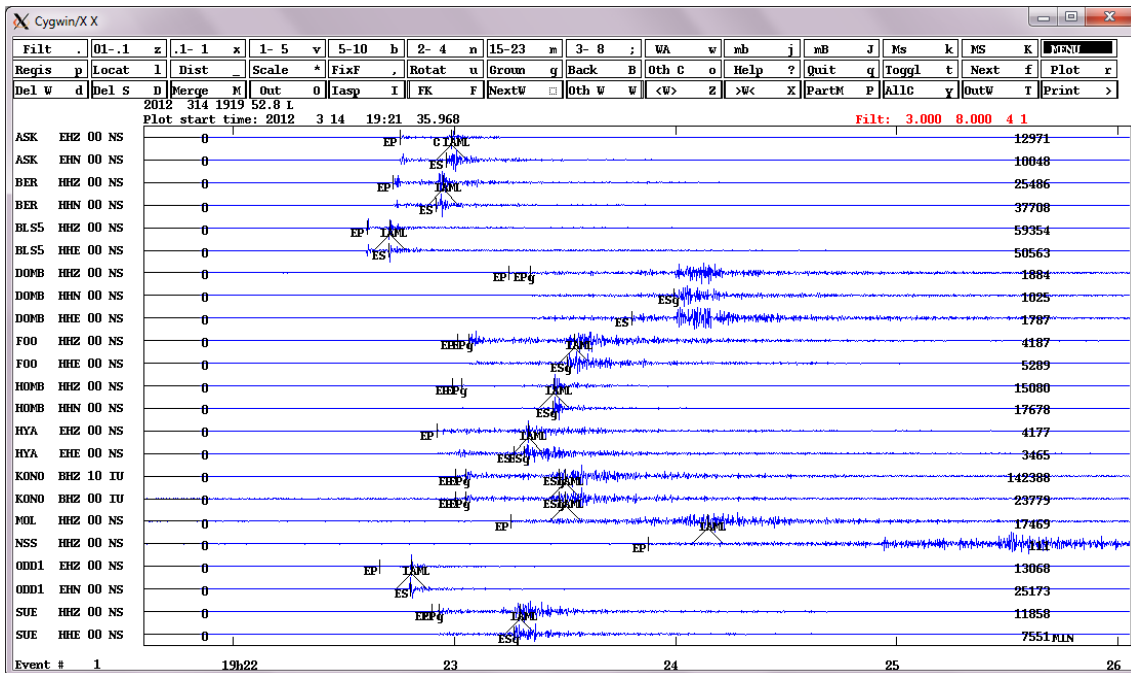
Reading different phases with different take-off angle from the source, such as Pg/Pn and Sg/Sn, significantly improves the depth determination. However, one should be quite certain that these identifications are correct. This requires fairly clear arrivals, an experienced analyst and that the phases match the solution when locating. For standard nomenclature, nature and travel paths of crustal phases see IS 2.1 and for many more record examples DS 11.1.



Three-component traces of station FOO at epicentral distance of 233 km, filtered 3-8 Hz: This is an example where Sn cannot be read, and S has to be read as Sg.



Single-trace plot of station NSS at epicentral distance of 649 km, filtered 5-10 Hz: Going to larger distances, the P arrival becomes more emergent and would not be read without applying a filter.



Multi-trace plot of all picked channels: At the end we look at all traces with phase onset and amplitude readings to see whether they are consistently picked. We then start locating. Note that there are simple rules of thumb for a first rough estimates of the approximate hypocenter distance, d , from identified first onsets of P and S waves. If these are P_g and S_g , then it holds $d[\text{in km}] = (S_g - P_g)[\text{in s}] \times 8$ and in the case of P_n and S_n $d = (S_n - P_n) \times 10$. More accurately, the multiplication factor depends on the average P and S wave crustal velocities [see eq. (3) in EX 11.1]: $d = t(S_g - P_g) (v_p v_s) / (v_p - v_s)$.

#	1	14	Mar	2012	19:22	28	L	59.526	5.615	13.9	0.4	3.4LBER	12	? 1
date	hrmn	sec	lat	long	depth	no	m	rms	damp	erln	erlt	erdp		
12	314	1922	28.19	5931.58N	5	36.9E	13.9	29	3	0.36	0.000	4.7	2.0	3.6
stn	dist	azm	ain	w	phas	calc	phs	hrmn	tsec	t-obs	t-cal	res	wt	di
BLS5	49	103.3	96.7	0	S	SG		1922	42.3	14.14	14.18	-0.04	1.00	7
BLS5	49	103.3	96.7	0	P	D PG		1922	36.4	8.15	8.15	0.00	1.00	4
BLS5	49	103.3		0	IAML			1922	42.5	14.3				
BLS5	49	103.3		0	AMPG			1922	36.4	8.2				
BLS5	49	103.3		0	AMSG			1922	42.4	14.2				
ODD1	71	52.6	93.0	0	S	SG		1922	48.0	19.81	20.02	-0.21	1.00	10
ODD1	71	52.6	93.0	0	P	D PG		1922	39.8	11.54	11.51	0.04	1.00	1
ODD1	71	52.6		0	IAML			1922	48.3	20.1				
BER	97	350.7	91.8	0	S	SG		1922	55.0	26.75	26.70	0.05	1.00	4
BER	97	350.7	91.8	0	P	C PG		1922	43.4	15.22	15.35	-0.12	1.00	2
BER	97	350.7		0	IAML			1922	56.9	28.7				
ASK	109	347.7	91.5	0	S	SG		1922	57.7	29.52	29.95	-0.43	0.98*	4
ASK	109	347.7	91.5	0	P	C PG		1922	45.3	17.06	17.21	-0.15	0.98*	3
ASK	109	347.7		0	IAML			1922	59.4	31.2				
SUE	177	344.9	90.8	0	Sg	SG		1923	16.2	48.02	47.82	0.20	0.87*	3
SUE	177	344.9	55.1	0	P	C PN4		1922	53.8	25.56	25.93	-0.37	0.87*	4
SUE	177	344.9	90.8	0	Pg	D PG		1922	55.8	27.61	27.49	0.13	0.87*	2
SUE	177	344.9		0	IAML			1923	17.7	49.5				
HYA	185	9.4	55.1	0	S	SN4		1923	16.1	47.86	46.91	0.95	0.86*	10
HYA	185	9.4	90.7	0	Sg	SG		1923	19.0	50.78	50.00	0.78	0.86*	3
HYA	185	9.4	55.1	0	P	C PN4		1922	55.3	27.11	26.96	0.15	0.86*	2
HYA	185	9.4		0	IAML			1923	20.1	51.9				
HYA	185	9.4	90.7	0	Pg	PG		1922	57.1	28.85	28.74	0.12	0.86*	0
HOMB	217	128.8	90.6	0	Sg	SG		1923	26.8	58.57	58.49	0.08	0.80*	6
HOMB	217	128.8	90.6	0	Pg	C PG		1923	2.1	33.86	33.62	0.25	0.80*	4
HOMB	217	128.8		0	IAML			1923	27.2	59.0				
HOMB	217	128.8	55.1	0	P	C PN4		1922	59.4	31.24	30.96	0.28	0.80*	6
KONO	225	84.8	55.1	0	P	PN4		1923	0.3	32.12	31.95	0.17	0.79*	3
KONO	225	84.8	90.6	0	Pg	D PG		1923	2.9	34.70	34.81	-0.11	0.79*	1
KONO	225	84.8	90.6	0	Sg	SG		1923	28.5	60.31	60.58	-0.26	0.79*	5
KONO	225	84.8		0	IAML			1923	30.2	62.0				
FOO	233	352.5	90.6	0	Sg	SG		1923	30.7	62.45	62.57	-0.12	0.78*	2
FOO	233	352.5	55.1	0	P	PN4		1923	0.9	32.71	32.88	-0.17	0.78*	2
FOO	233	352.5	90.6	0	Pg	PG		1923	3.9	35.69	35.96	-0.27	0.78*	1
FOO	233	352.5		0	IAML			1923	32.5	64.3				
DOMB	342	32.3	55.1	0	S	SN4		1923	48.1	79.93	80.82	-0.89	0.60*	6
DOMB	342	32.3	90.4	0	Sg	SG		1923	59.5	91.25	91.28	-0.03	0.60*	3
DOMB	342	32.3	55.1	0	P	C PN4		1923	14.9	46.70	46.45	0.26	0.60*	1
DOMB	342	32.3	90.4	0	Pg	PG		1923	20.6	52.35	52.46	-0.11	0.60*	0
MOL	355	16.3	55.1	0	P	PN4		1923	15.4	47.23	48.03	-0.80	0.57*	1
MOL	355	16.3		0	IAML			1924	8.9	100.7				
NSS	649	28.0	53.1	0	P	PN5		1923	52.6	84.43	83.82	0.61	0.08*	0
ODD1	EZ	hdist:	72.6	amp:	4529.1	T:	0.4	ml =	3.7					
BER	HZ	hdist:	97.8	amp:	1863.0	T:	0.4	ml =	3.5					
ASK	EZ	hdist:	109.9	amp:	2311.9	T:	0.1	ml =	3.6					
SUE	HZ	hdist:	177.5	amp:	1124.8	T:	0.1	ml =	3.6					
HYA	EZ	hdist:	185.5	amp:	629.6	T:	0.4	ml =	3.4					
HOMB	HZ	hdist:	218.4	amp:	1355.1	T:	0.4	ml =	3.8					
KONO	BZ	hdist:	225.4	amp:	144.8	T:	0.2	ml =	2.8					
MOL	HZ	hdist:	355.3	amp:	112.3	T:	0.1	ml =	3.0					
FOO	HZ	hdist:	233.4	amp:	388.3	T:	0.4	ml =	3.3					
BLS5	HZ	hdist:	50.9	amp:	6142.4	T:	0.1	ml =	3.7					
BLS5	HZ	gdist:	50.9	mom:	13.5	mw =	2.9							
FOO	HZ	gdist:	231.4	mom:	13.6	mw =	3.0							
Number of spectra available and number used in average														
2012	314	1922	28.2	L	59.526	5.615	13.9	BER	12	0.4	3.4LBER	3.0	WBER	

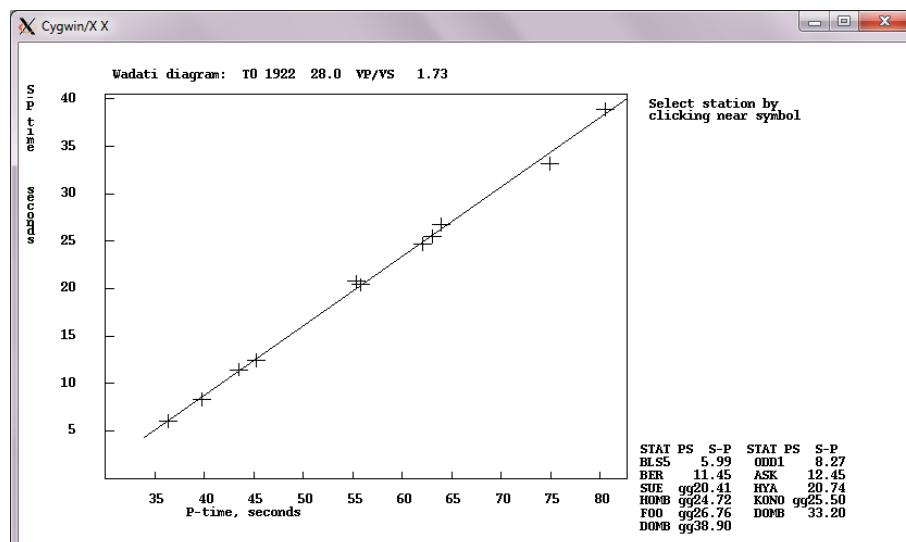
Output from location program: We use a standard linearized inversion program (HYPOCENTER, Lienert et al. (1986, 1995)) to locate the event that makes use of the common seismic phases at local distances. The location program tries to minimize the differences between the observed (t-obs) and calculated arrival times (t-cal) that are computed for a 1D layered model. The residuals are given in the column labeled 'res'. For our example event these travel-time residuals are relatively small. Note that for station DOMB the residuals are given for both Pn/Pg and Sn/Sg and that our interpretation fits the solution.

Direct (Pg and Sg) and critically refracted wave types (Pn and Sn) were very clearly visible from this earthquake on the records of several stations. This helped to constrain the earthquake depth, which was estimated to be 13.9 km (see last but one line). At the end the location procedure the program gives the magnitude values for the individual components and stations, and calculates the average, which is 3.4 in this case. The text 'LBER' refers to the magnitude type (L for ML) and the reporting agency (BER = Bergen). Some spread in the magnitude residuals (difference between single-component magnitude and average) is expected and mostly due to local site effects. In our case the largest residual are -0.6 and +0.4 magnitude units.

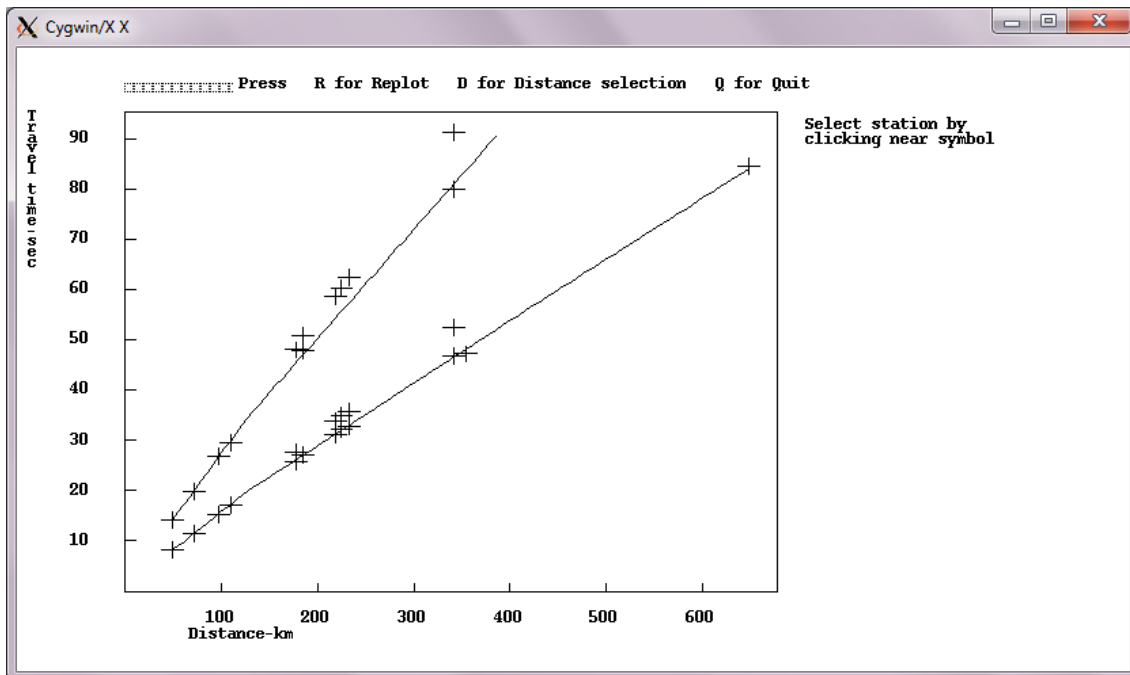
A number of polarities were read, 7 compressional (C) and 4 dilatational (D), which are given in the column after the phase name read and the phase used by the location program. For example, in the second phase line of the location output for stations BLS5, the first phase arrival was read as P, the location program decides that this is the direct Pg arrival based on the hypocentral distance. The polarity is read as D, standing for down or dilatation. The location program computes the azimuth (azm) and angle of incidence (ain), which in this example are 103.3° and 96.7°, respectively. The polarity data can be used to determine the fault plane solution if there are sufficient observations. In this example we can not find a robust solution and do not show the procedure here.

The velocity model used is given by the following table:

<i>Depth of top of layer in km</i>	<i>P-wave velocity in km/s</i>
0	6.2
12	6.6
23	7.1
31	8.05
50	8.25
80	8.5

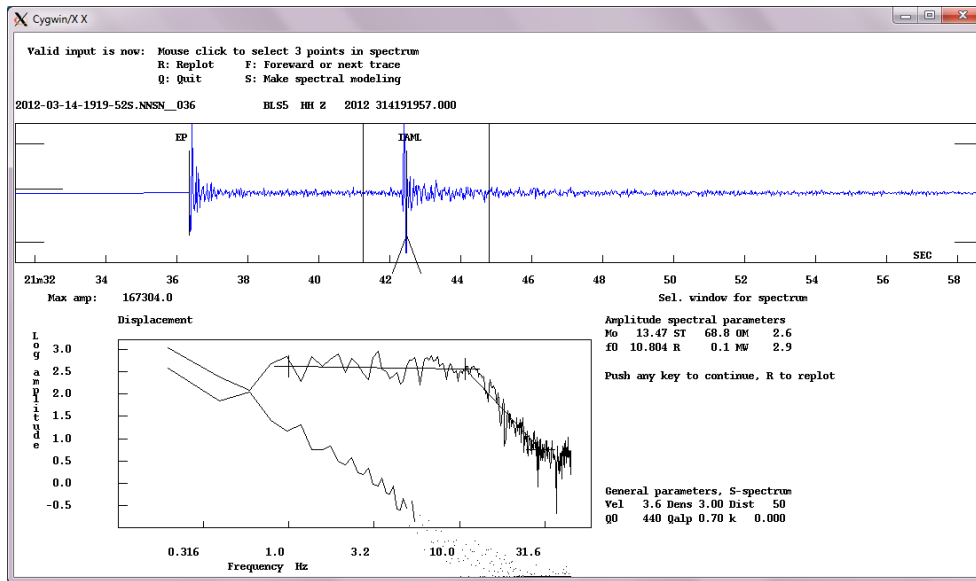


Wadati diagram: To check our readings we can plot the Wadati diagram (plotting S-P times against absolute P arrival times (labels on x-axis given to show scaling), and the data points should fall onto a line like in this example. Data points (+) that are significantly off the line should be checked and re-read.

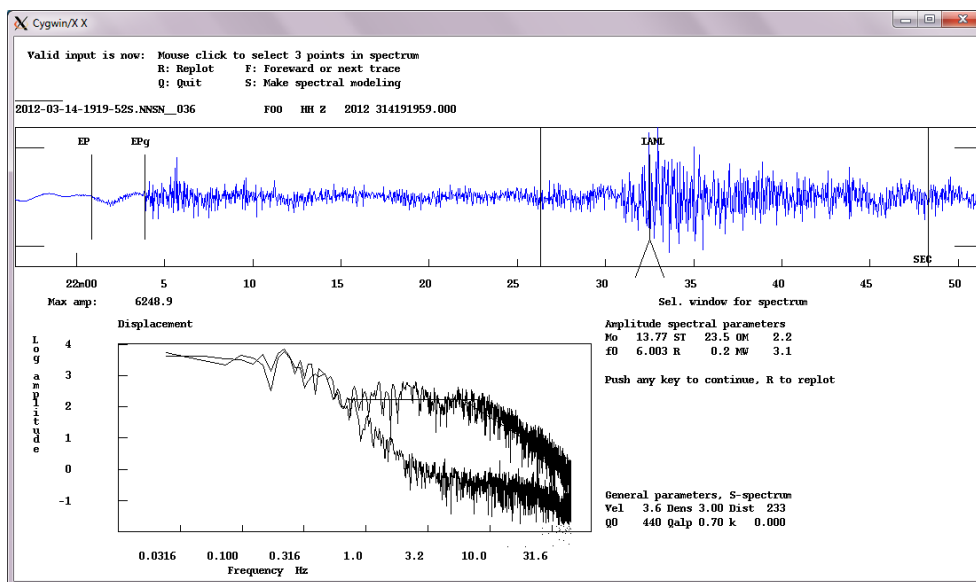


Travel time plot: We can also produce a simple travel-time plot that shows the observations (+) together with the calculated travel-time curves for P and S first arrivals according to the local velocity model implemented in the SEISAN location program (solid lines). For distances greater than about 150 km, our Pg (around the lower curve) and Sg arrivals (around the upper curve) come in mostly after the predicted first arrivals.

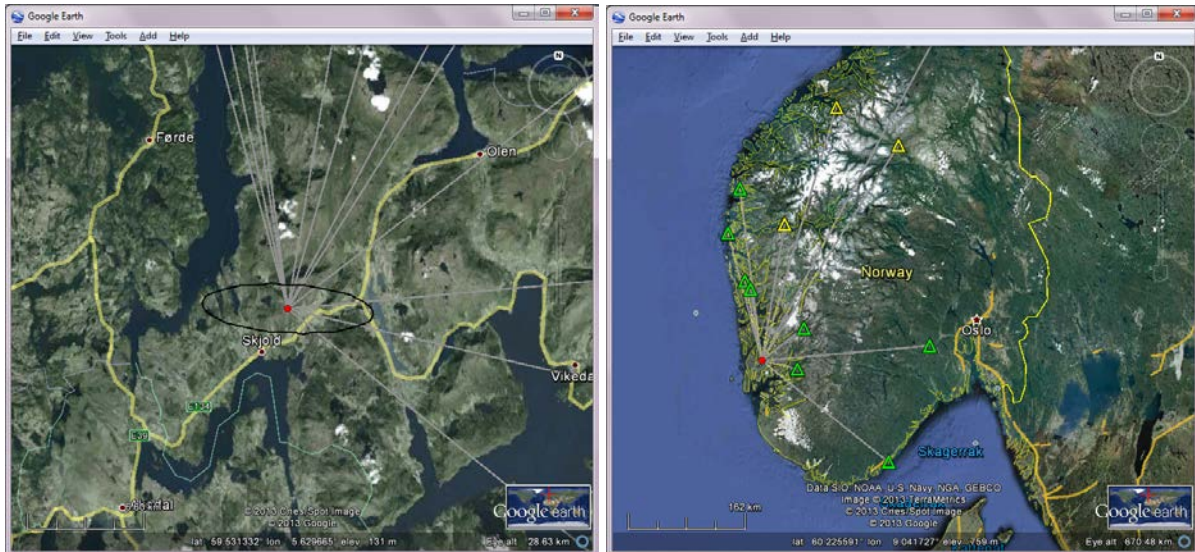
Note that SEISAN uses layered 1D velocity models and the user needs to enter the model for his or her respective area. SEISAN allows the use of a number of 1D models for events depending on their location to account for changes in structure within the area of interest or network coverage.



Source spectrum plot for station BLS5 at epicentral distance of 49 km: We compute the source displacement spectrum (upper curve in the bottom diagram above the noise spectrum) for the vertical component of this station. The spectrum is corrected for attenuation, which for Norway is given by $Q(f) = 440 \times f^{0.7}$ (see data at bottom right). From the spectrum we can approximately read the long-period spectral level, which is translated via the assumed velocity and density model into seismic moment (here $\log_{10}M_0 = 13.47$ Nm) and moment magnitude ($M_W = 2.9$). We also measure the corner frequency ($f_c = 10.8$ Hz), which is translated into stress drop $\Delta\sigma = 68.8$ bars = 6.88 MPa using $\Delta\sigma = (7/16) M_0/R^3$ (Eshelby, 1957).



Source spectrum plot for station FOO at epicentral distance of 233 km: The spectral analysis is normally done for a number of stations with good SNR. When comparing the results for this station with those of the previous station (BLS5) one finds the difference between the moment magnitudes to be small, but the stress drop difference to be large because of the reading uncertainty in f_c and $\Delta\sigma \sim f_c^3$. (For related formulas and discussions see EX 3.4). At the end, the averages are computed from all observations and standard deviations are calculated.



Location maps in GoogleEarth: SEISAN creates files that make it possible to look at the location with GoogleEarth immediately. The figure on the left shows the epicenter area with the error ellipse plotted as a solid black line. The extent of the 90 percent confidence error ellipse is about 2.5 by 9.0 km. It is also possible to compare the epicenter location to known fault structures. In the case of explosions we try to see if the event is located close to a known mine or quarry.

On the right, we show a more regional GoogleEarth map, which shows the epicenter as a red dot. The stations are color coded depending on their residual from the hypocenter location where green is low residual, yellow is moderate residual and red is high residual (the color levels are configurable). In this example, there are no stations towards the west and southwest of the epicenter and stations are sparse toward the southeast, resulting in an azimuthal gap of 216 degree. This, together with the dominating N-S extension of the Norwegian network causes the much larger location uncertainty in E-W direction (see error ellipse in the left-hand panel). We can also plot the epicenter together with previously located earthquakes.

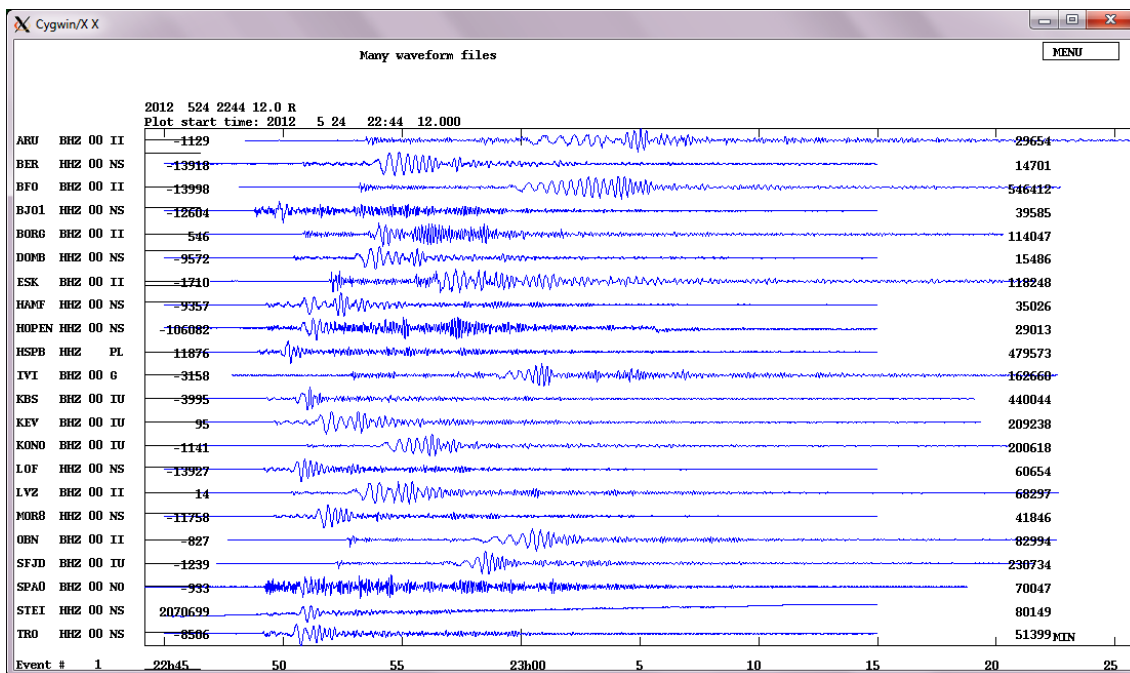
3 Analysis of a regional earthquake recorded at distance between 440 and 2950 km ($\Delta < 30^\circ$)

As an example we have chosen an earthquake that occurred on the Mohns Ridge, offshore northern Norway. The earthquake with a magnitude greater than 6 was large enough to be recorded globally. Here, we process data that were recorded on the Norwegian National Seismic Network operated by the University of Bergen and data provided by IRIS from the Global Seismograph Network (GSN). All data used here was recorded on broadband stations.

The earthquake parameters provided by the USGS are:

Date:	24 May 2012
Origin time:	22:47:46.6 (UTC)
Latitude:	72.99°N
Longitude:	5.65°E
Depth:	8.8 km
Mw:	6.2

We now show the step-by-step procedure to analyze the earthquake.

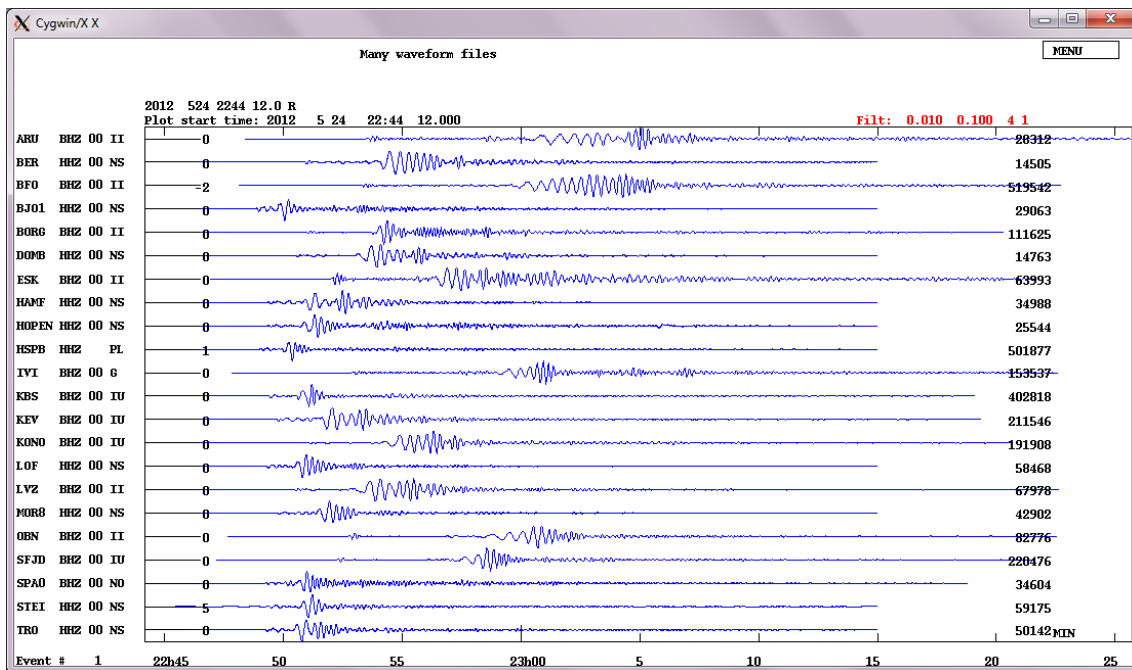


Vertical component multi-trace plot, unfiltered: On the seismograms we clearly see long-period surface waves, with large amplitudes and pronounced dispersion. This indicates that the earthquake is relatively large and shallow. For the stations, where we see the first arriving P as well we can estimate the approximate distance from the time difference between the P and the Rayleigh surface wave maximum (which is most pronounced in vertical component records). For station BFO this time difference is about 10 min. According to Table 1 below this means that station BFO is about 20 to 25° degrees away ($1^\circ = 111.2$ km) from the epicenter. Being aware of the station locations, considering the pattern of arrival times, and,

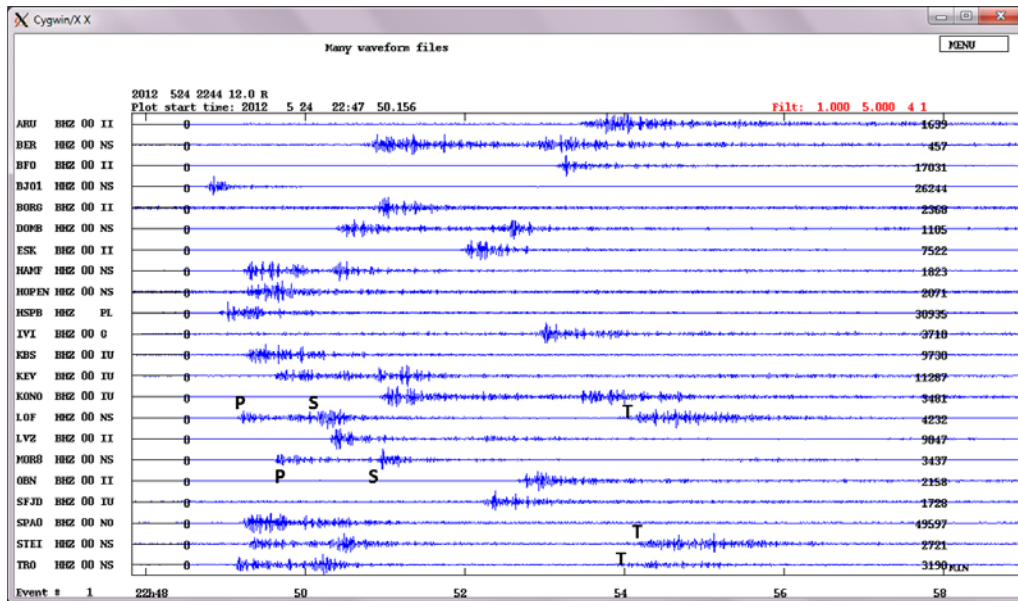
if three-component records are available, also of the first-motion polarity patterns and the better constraining S-P travel-time differences, we can also get a first guess of the epicenter location (see EX 11.2).

Table Approximate time interval ($t_{Rmax} - t_p$) between the arrival of the maximum Rayleigh wave amplitude and the first onset of P waves as a function of epicentral distance Δ according to Archangelskaya (1959) and Gorbunova and Kondorskaya (1977) for dominantly continental travel paths.

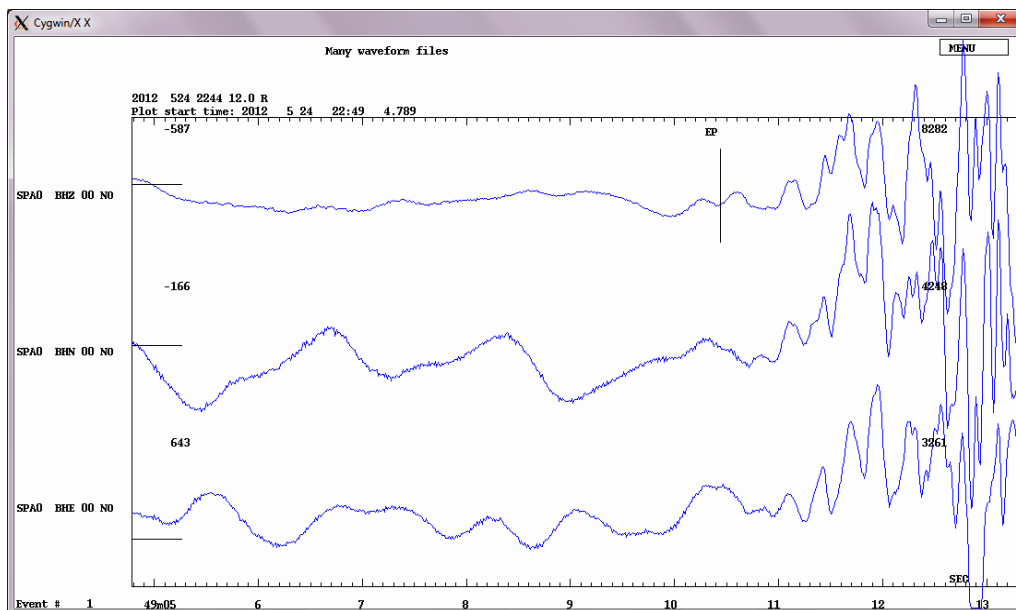
Δ°	$t_{Rmax} - t_p$ (min)	Δ°	$t_{Rmax} - t_p$ (min)	Δ°	$t_{Rmax} - t_p$ (min)
10	4-5	55	26	100	45-46
15	6-8	60	28-29	105	47-48
20	9-10	65	31	110	48-50
25	10-12	70	33	115	53
30	13-14	75	35	120	55
35	15-16	80	37	125	57
40	18-19	85	39-40	130	60
45	21	90	42	140	64
50	24	95	43	150	70



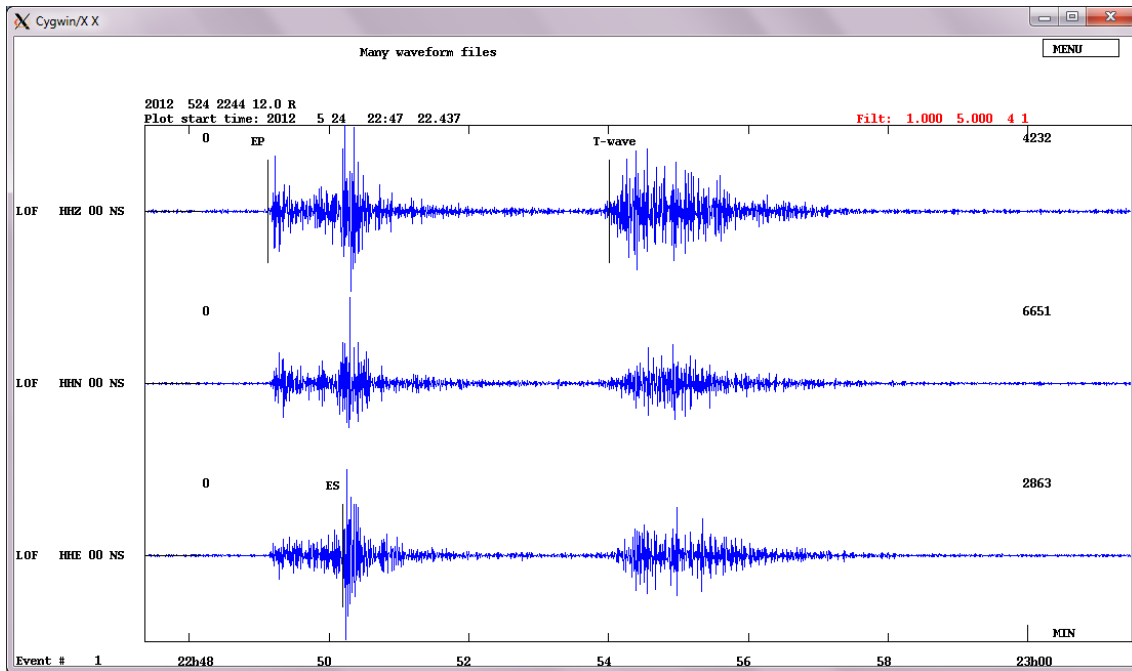
Vertical component multi-trace plot, filtered 0.01-0.1 Hz: Filtering the data at low frequencies, we see both body and surface waves. However, the seismograms are dominated by the surface waves as they have larger amplitudes in this frequency range. The shape of the surface wave trains depends on the Earth structure along the paths (see, e.g., Chapter 2, Figs. 2.10, 2.12 and 2.13) and, because of the dispersion, i.e., the dependence of the surface-wave velocity on the period, also on the epicentral distance (see Figure above). The seismograms also depend on the mechanism of faulting.



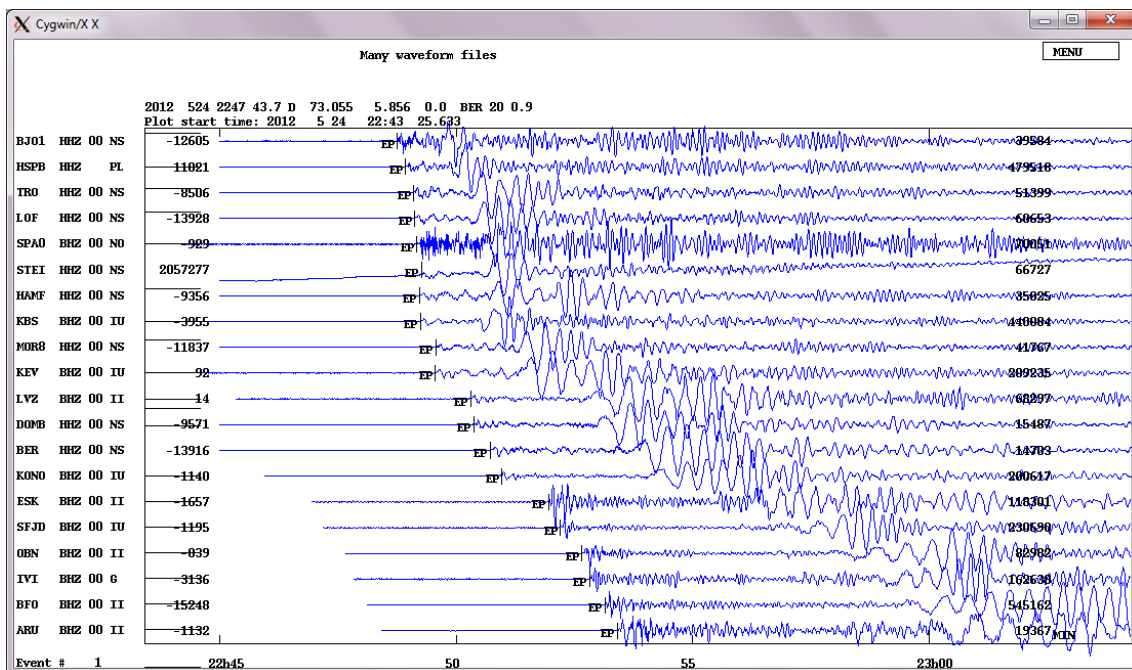
Multi-trace plot, filtered 1-5 Hz: Looking at higher frequencies, the surface waves disappear from the seismogram and we look at P and (at these high frequencies and in vertical component records usually less clear) S body waves. If the first arriving P and S phases are P_n and S_n, respectively, then there is a simple “rule of thumb” to estimate the hypocenter distance, namely time difference S_n-P_n [in s] × 10. For station LOF with a time difference of about 1 min this means a distance of ≈ 600 km, This agrees with the rough estimate from the time difference between the Rayleigh-wave maximum and the P first arrival. On some stations (STEI, TRO and LOF) we can also see a signal in this frequency range arriving much later (about 5 minutes after P for station TRO). This is the T(tertiary)-wave, generated by an earthquake in or near the oceans, propagating in the oceans as an acoustic wave and converted back to seismic wave at the ocean-land boundary near the recording site. (for more explanations see Chapter 2, section 2.6.5). It should not be mistaken for another event.



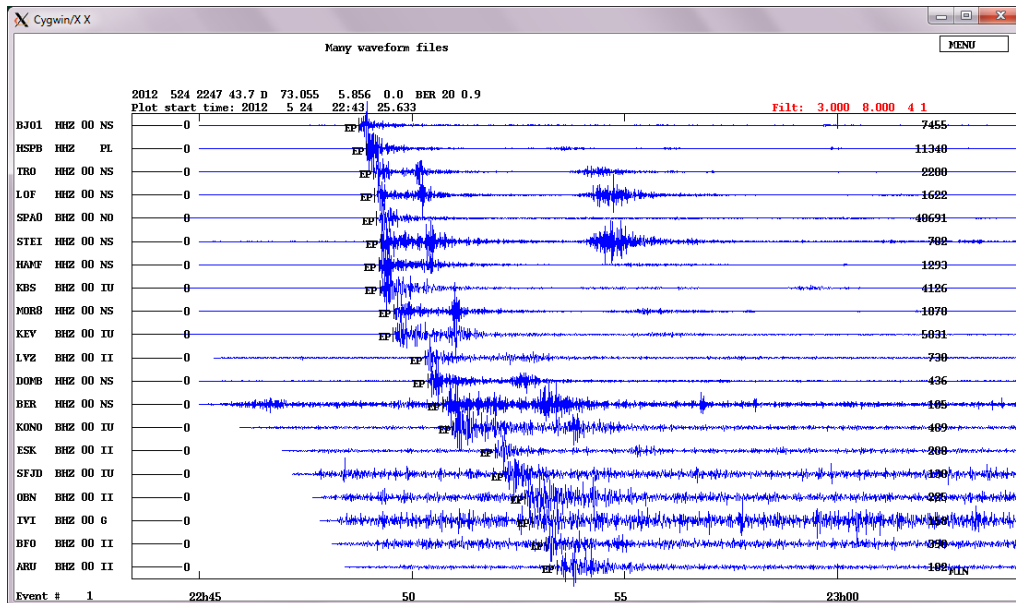
Three-component traces of station SPA0 at epicentral distance of 648 km, unfiltered: Zooming around the P arrival the phase can be picked with or without filter. The error on the phase reading here may be as large as 0.5 sec.



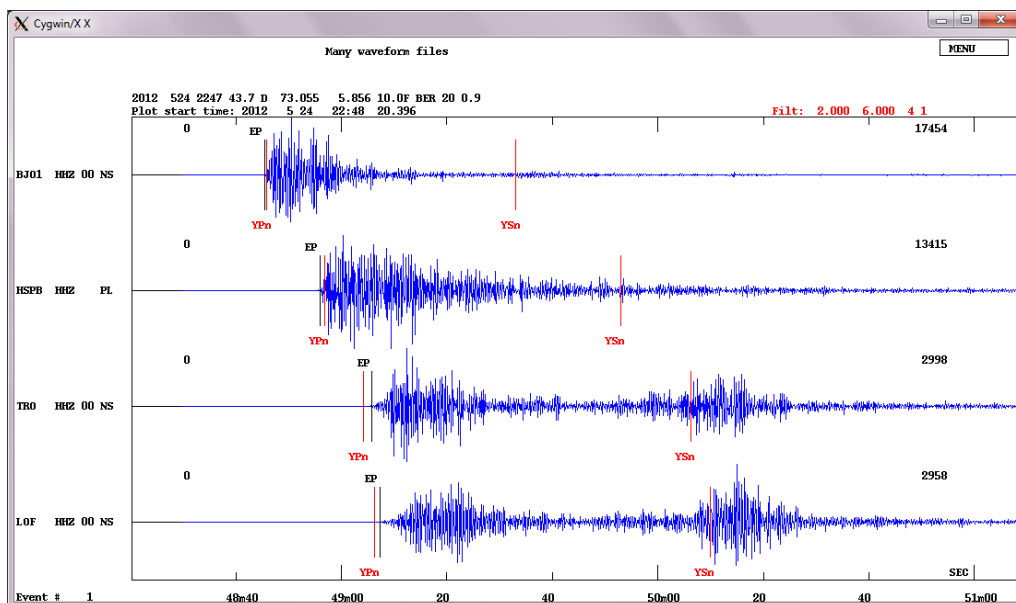
Three-component traces of station LOF at epicentral distance of 603 km, filtered 1-5 Hz: Looking at this coastal station, we clearly see the arrival of the T wave. The amplitude of the T wave is in this case comparable to the S-wave amplitude. According to the time difference (S-P)[in s]×10 the station is about 600 km away from the source.



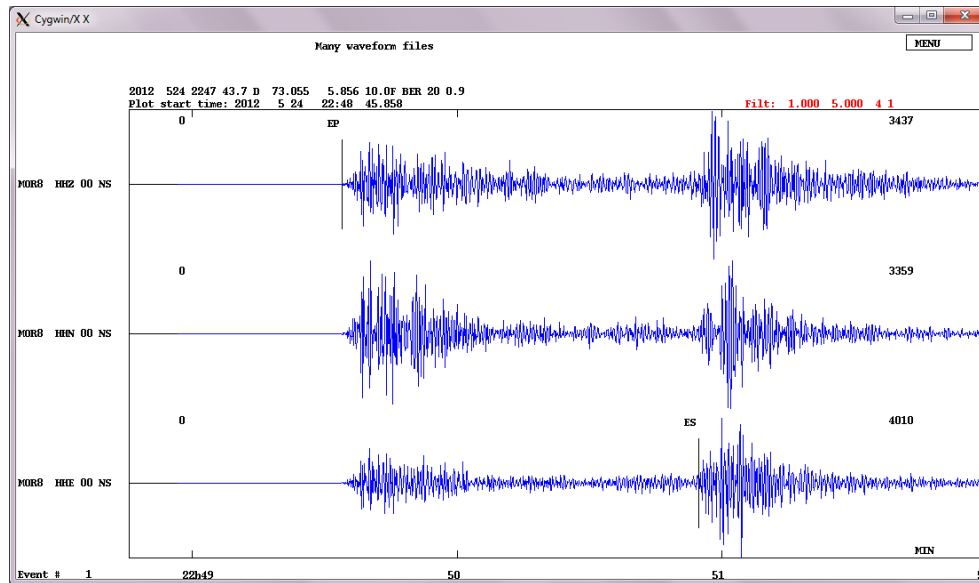
All vertical traces, unfiltered: The traces are sorted by distances and the P arrival picks are shown. This type of plot is useful to check the consistency of the picks. Based on the P arrival times, the epicenter is estimated at 73.055°N and 5.856°E. The hypocenter depth is not constrained by the data and given as 0 km by the location program. On stations like SPA0 we can see that we have both high and low frequency. Towards larger distances the high frequencies get attenuated more strongly than lower frequency signals.



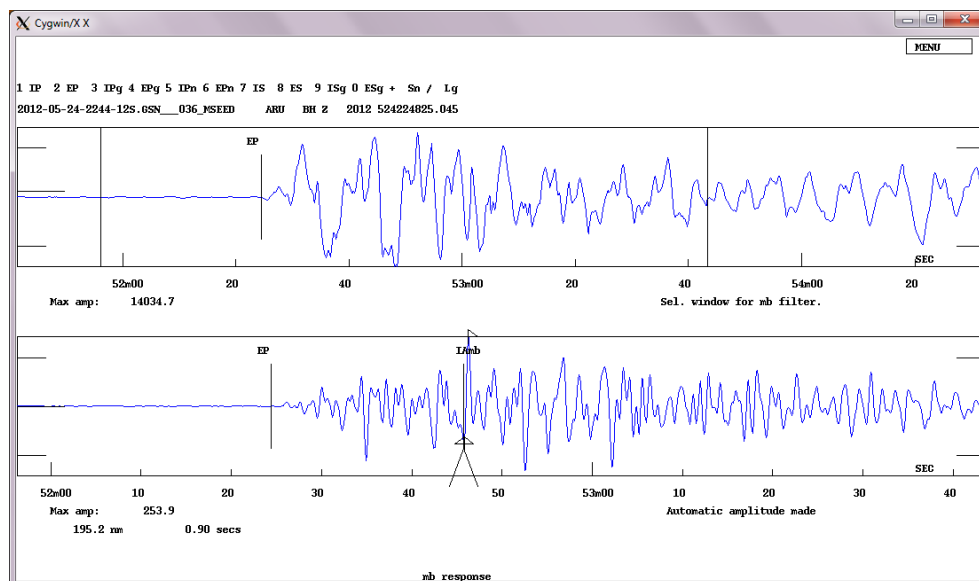
All vertical traces, filtered 3-8 Hz: We can see that the signal amplitude of the higher frequencies decreases as we go to larger distances due to attenuation. The distance range here is approximately 440 km (station BJO1) to 2960 km (station ARU).



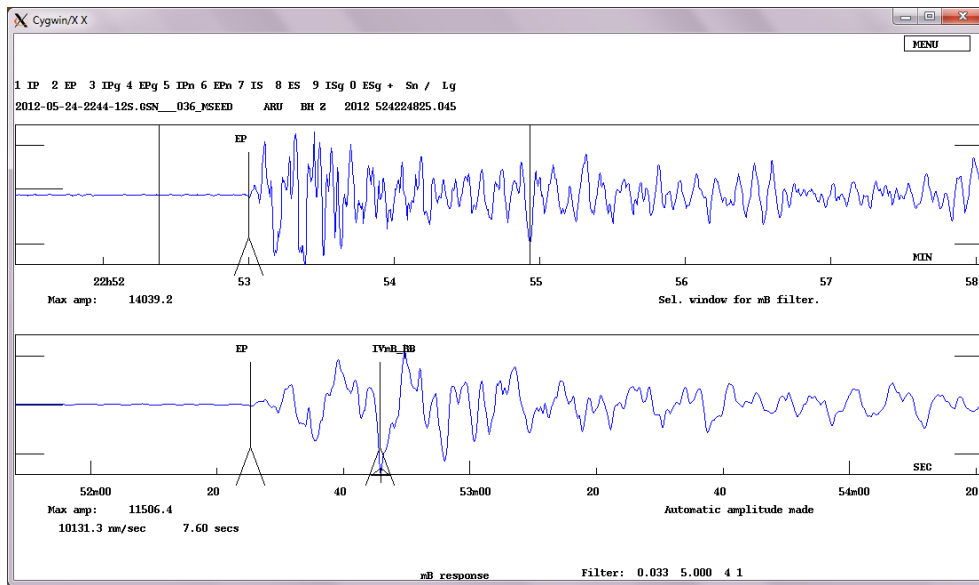
Selected vertical traces, filtered 2-6 Hz: Computed Pn and Sn phase arrival times have been marked by red vertical lines and the text 'yPn' and 'ySn'. No S-waves are recognizable in the records of the two stations BJO1 and HSPB closer to the epicenter. They are likely situated in an azimuth range of minimum S-wave and strong P-wave radiation (see EX 3.2). While S appears at larger distances, the phase is emergent and from this plot its travel-time does not match well with the global velocity model used (IASP91). The earthquake is located on the Mid-Atlantic Ridge. The seismograms are complicated both due to source and path effects, including the transition from oceanic to continental crust. The display of theoretical arrivals should only serve as a guide on whether the considered phase is plausible to arrive in this time window. However, it would be **wrong** to pick a phase close to the predicted time if there is no visible change in the signal.



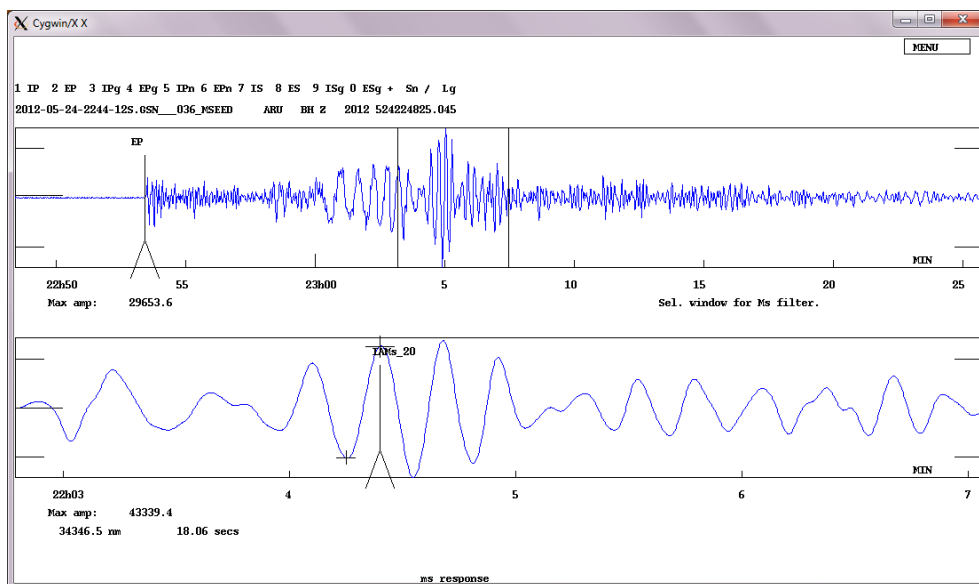
Three-component trace plot of station MOR8 at epicentral distance of 813 km, filtered 1-5 Hz: This is an example where the S arrival can be read quite well on one of the horizontal channels.



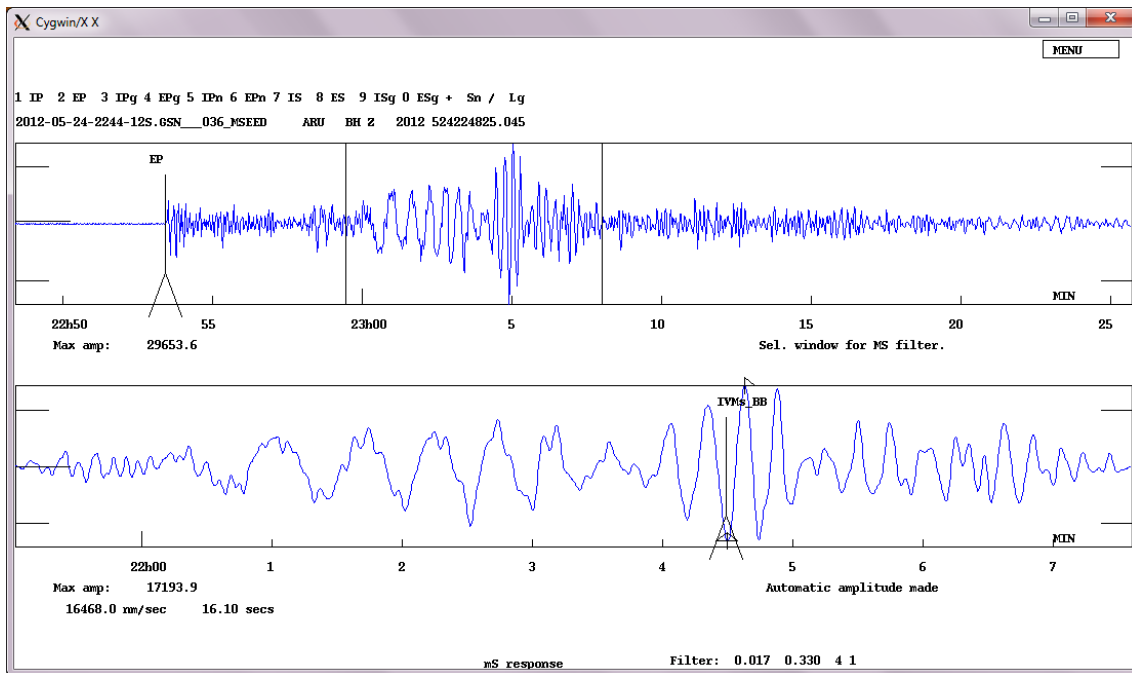
Vertical-component trace of station ARU ($\Delta=26.7^\circ$), mb filter: While it is possible to read ML amplitudes from this earthquake for stations within about 1500 km, this is not done because the ML scale for Norway is not appropriate for sources on the Mid-Atlantic Ridge. We, therefore, read amplitudes for the teleseismic magnitudes which can be used down to distances of 20° and even down to 2° for standard M_s_BB . Moment magnitude for this event can be measured using either regional or teleseismic moment tensor inversion (USGS $M_w = 6.2$). The plot above shows the mb amplitude reading for station ARU. This is done by deconvolving the broadband instrument response and applying a filter to simulate the WWSSN short-period seismograph (see IS 3.3). The simulation trace is displayed on the bottom trace for the time window indicated by the vertical lines on the top trace (raw data). The amplitude Iamb is automatically read between the largest peak and adjacent trough. It is stored as half of this measurement, which is 195.2 nm, together with the related period T of 0.90 s which is twice the time difference between peak and trough.



Vertical-component trace of station ARU ($\Delta=26.7^\circ$), mB_BB filter: This plot shows the mB_BB amplitude reading made for station ARU. This is done by deconvolving the actual instrument response to a broadband velocity response and reading the velocity amplitude on the simulation trace which is displayed on the bottom for the time window indicated by the vertical lines on the top trace (raw data). The amplitude is automatically read between peak and trough, and stored as half of this measurement, which is 10131.3 nm/s with a period of 7.60 s. The period for mB_BB needs to be within $0.2 \text{ s} < T < 30 \text{ s}$ (see IS 3.3).



Vertical-component trace of station ARU ($\Delta=26.7^\circ$), Ms_20 filter: This plot shows the Ms_20 amplitude reading made for station ARU. This is done by deconvolving the displacement instrument response and applying a filter to simulate the WWSSN long-period seismograph. The simulation trace is displayed on the bottom for the time window indicated by the vertical lines on the top trace (raw data). The amplitude is automatically read between peak and trough, and stored as half of this measurement, which is 34346.5 nm with a period of 18.06 s. The period for Ms_20 needs to be within $18 \text{ s} \leq T \leq 22 \text{ s}$.

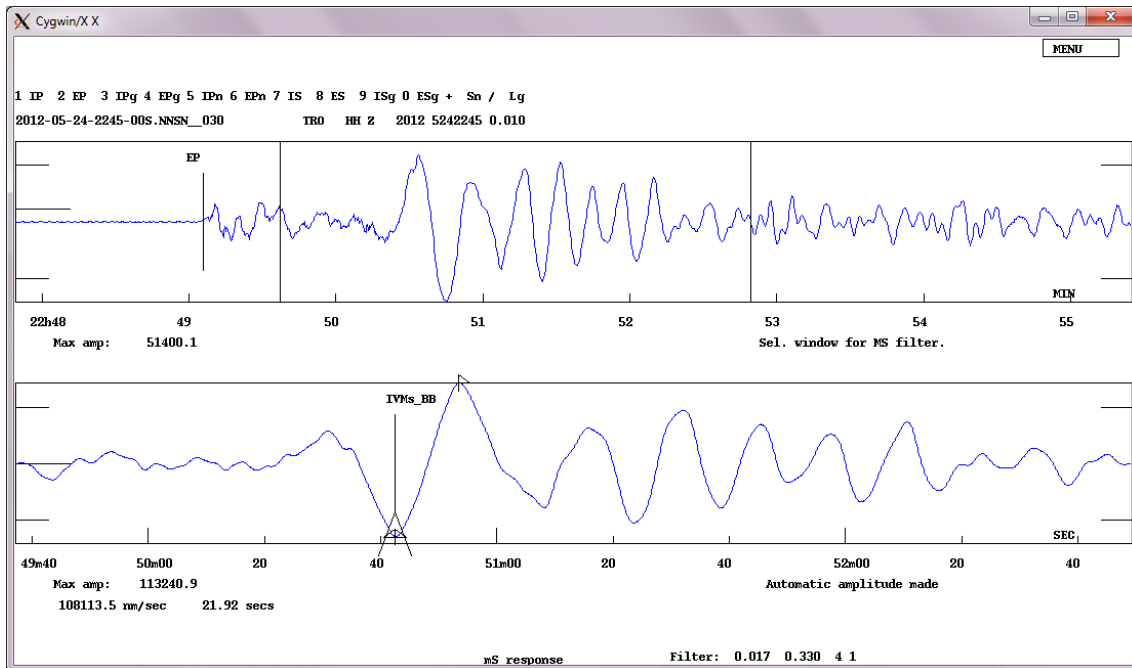


Vertical-component trace of station ARU ($\Delta = 26.7^\circ$), *Ms_BB* filter: This plot shows the *Ms_BB* amplitude reading made for station ARU. If the primary record is not yet a velocity broadband trace then the latter has to be synthesized by deconvolving the actual instrument response to a broadband velocity response and reading the velocity amplitude on the simulation trace. The latter is displayed on the bottom trace for the time window indicated by the vertical lines on the top trace (raw data). The amplitude is automatically read between the largest peak and adjacent trough, and stored as half of this measurement, which is 16468.0 nm/s, together with a period of 16.10 s, i.e., twice the time difference between the peak and adjacent trough. The period needs to be within $3 \text{ s} \leq T \leq 60 \text{ s}$, and thus maybe significantly shorter than for *Ms_20* measurement, allowing *Ms* measurements at distances well below 20° (down to 2°).

Based on these four amplitude and period measurements from station ARU, we can compare the magnitudes to *Mw*:

Magnitude scale	Value
<i>Mw</i> (USGS)	6.2
<i>mb</i>	5.9
<i>mB_BB</i>	6.7
<i>Ms_20</i>	6.0
<i>Ms_BB</i>	6.1

We see that the values for this size of earthquake around magnitude 6 compare well, with the exception of *mB_BB* which is on global average about 0.3 -0.4 m.u. greater than *Mw* at *Mw* = 6.2 and about 0.6 m.u larger than *mb*.



Vertical-component trace of station TRO at $\Delta=5.3^\circ$, M_s_{BB} filter: M_s_{BB} can be used down to distances of 2° . This plot shows the M_s_{BB} amplitude reading made for station TRO. The amplitude is automatically read between peak and trough, and stored as half of this measurement, which is 108113.5 nm/s with a period of 21.92 s. The computed magnitude is $M_s_{BB} = 5.7$, compared to $M_s_{BB} = 6.1$ for station ARU.


```

# 1 24 May 2012 22:47 45 D 72.943 6.096 8.8F 1.8 6.3sBER 20 ? 1

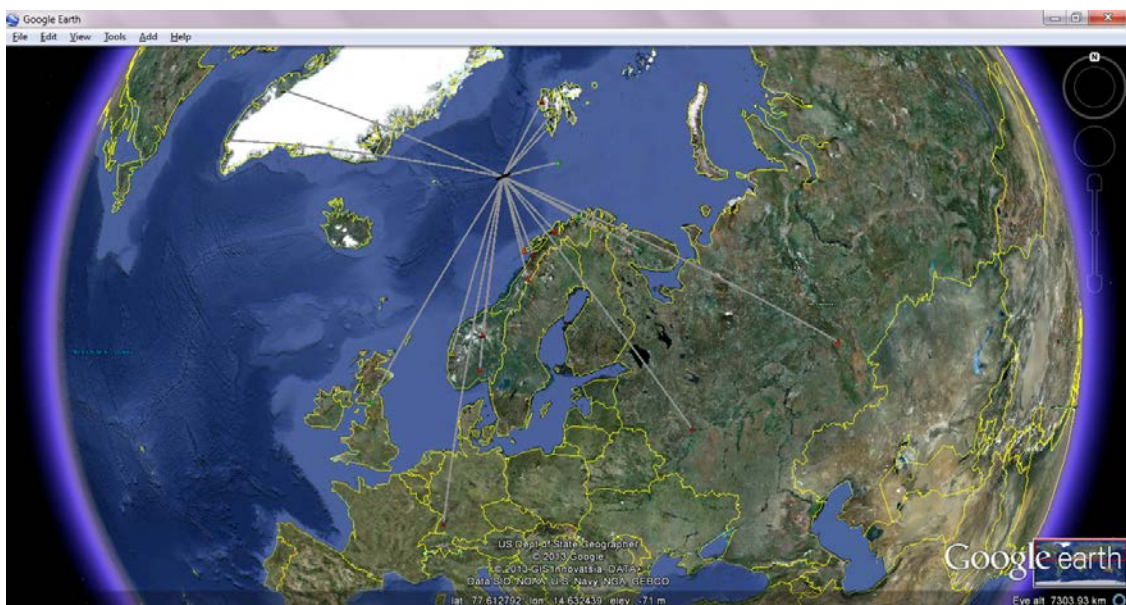
date hrmn sec lat long depth no m rms damp erln erlt erdp
12 524 2247 45.09 7256.61N 6 5.7E 8.8* 24 2 1.77 0.000 47.0 9.2 0.0
stn dist azm ain w phas calcphs hrmn tsec t-obs t-cal res wt di
BJO1 439 60.6 45.8 0 P Pn 2248 45.4 60.36 60.82 -0.46 1.00 6
HSPB 528 26.6 45.8 0 P Pn 2248 56.0 70.88 71.85 -0.97 1.00 7
TRO 587 122.6 45.8 0 P Pn 2249 5.8 80.70 79.19 1.51 1.00 1
TRO 587 122.6 0 IVMs_BB 2250 42.4 177.3
LOF 603 149.2 45.8 0 P Pn 2249 7.4 82.27 81.18 1.09 1.00 1
LOF 603 149.2 48.3 0 S Sn 2250 6.4 141.35 144.88 -3.53 1.00 5
SPA0 648 21.2 45.8 0 P Pn 2249 10.4 85.35 86.84 -1.49 1.00 7
STEI 654 144.2 45.8 0 P Pn 2249 15.7 90.63 87.49 3.14 1.00 1
STEI 654 144.2 48.2 0 S Sn 2250 18.9 153.78 156.21 -2.43 1.00 6
HAMF 662 104.2 45.8 0 P Pn 2249 13.8 88.70 88.50 0.20 1.00 3
KBS 685 10.5 45.8 0 P Pn 2249 14.6 89.50 91.27 -1.77 1.00 8
MOR8 813 151.6 45.7 0 P Pn 2249 34.0 108.87 107.24 1.63 1.00 1
MOR8 813 151.6 48.1 0 S Sn 2250 54.7 189.65 191.69 -2.04 1.00 5
KEV 821 105.3 45.7 0 P Pn 2249 33.8 108.76 108.20 0.56 1.00 3
LVZ 1191 103.8 45.6 0 P Pn 2250 18.8 153.67 153.99 -0.32 1.00 3
DOMB 1219 172.5 45.6 0 P Pn 2250 22.6 157.46 157.34 0.12 1.00 2
DOMB 1219 172.5 47.8 0 S Sn 2252 25.0 279.88 281.53 -1.65 1.00 7
BER 1401 181.7 45.5 0 P Pn 2250 43.7 178.61 179.70 -1.09 1.00 3
KONO 1490 172.3 45.4 0 P Pn 2250 58.2 193.13 190.69 2.44 1.00 2
ESK 2010 197.3 39.8 0 P Pn 2251 57.3 252.18 251.84 0.34 1.00 5
ESK 2010 197.3 0 IVmB_BB 2252 6.0 260.9
ESK 2010 197.3 0 IAMB 2252 12.8 267.7
ESK 2010 197.3 0 IAMs_20 2257 17.9 572.8
SFJD 2172 281.2 34.8 0 P P 2252 11.8 266.68 267.71 -1.03 1.00 12
SFJD 2172 281.2 0 IAMB 2252 23.3 278.2
SFJD 2172 281.2 0 IVMs_BB 2258 30.4 645.3
SFJD 2172 281.2 0 IAMs_20 2258 34.0 648.9
OBN 2420 128.4 34.0 0 P P 2252 39.3 294.24 291.96 2.28 1.00 0
OBN 2420 128.4 0 IVmB_BB 2252 50.0 304.9
OBN 2420 128.4 0 IAMB 2252 55.4 310.3
OBN 2420 128.4 0 IAMs_20 23 0 30.7 765.6
OBN 2420 128.4 0 IVMs_BB 23 0 39.0 773.9
IVI 2570 268.3 33.4 0 P P 2252 50.1 304.96 306.21 -1.25 1.00 11
BFO 2743 176.4 28.4 0 P P 2253 9.1 324.01 321.15 2.86 1.00 1
ARU 2954 100.0 28.1 0 P P 2253 25.2 340.13 338.25 1.87 1.00 1
ARU 2954 100.0 0 IAMB 2253 45.7 360.6
ARU 2954 100.0 0 IVmB_BB 2253 45.8 360.7
ARU 2954 100.0 0 IAMs_20 23 4 49.9 1024.8
ARU 2954 100.0 0 IVMs_BB 23 4 54.1 1029.0
TRO HZ dist: 587.0 amp: 108000.0 T: 21.9 MS = 5.7
ESK BZ dist: 2010.0 amp: 154000.0 T: 7.1 mB = 7.3
ESK BZ dist: 2010.0 amp: 11890.7 T: 1.8 mb = 6.7
ESK BZ dist: 2010.0 amp: 363000.0 T: 21.6 Ms = 6.6
SFJD BZ dist: 2172.0 amp: 4562.0 T: 2.9 mb = 6.2
SFJD BZ dist: 2172.0 amp: 158000.0 T: 18.0 Ms = 6.4
OBN BZ dist: 2420.0 amp: 31652.9 T: 5.2 mB = 6.9
OBN BZ dist: 2420.0 amp: 6234.6 T: 2.4 mb = 6.6
OBN BZ dist: 2420.0 amp: 168000.0 T: 21.8 Ms = 6.4
OBN BZ dist: 2420.0 amp: 49453.8 T: 19.6 MS = 6.4
ARU BZ dist: 2954.0 amp: 269.7 T: 1.0 mb = 5.9
ARU BZ dist: 2954.0 amp: 10130.1 T: 7.6 mB = 6.7
ARU BZ dist: 2954.0 amp: 36322.6 T: 18.1 Ms = 6.0
ARU BZ dist: 2954.0 amp: 16468.0 T: 16.1 MS = 6.1
2012 524 2247 45.1 D 72.943 6.096 8.8F BER 20 1.8 6.3sBER 6.1sBER 6.4bBER
    
```

Output from location program: We use the same standard linearized inversion program (HYPOCENTER) to locate this event as used for the local event. The location is based on the IASP91 global travel-time model. The residuals are much larger than for the local event as the difference between the true regional model and IASP91 is significant. The largest residual here is more than 3 seconds. In the hypocenter determination we fixed the depth to 8.8 km and, using both P and S readings, we obtain a latitude of 72.943°N and longitude of 6.096°E. This compares to 73.055°N and 5.856°E for the preliminary location based on P arrivals only. The solution based on P and S is 15 km east of the USGS solution.

Based on all our amplitude and period measurements the following network average magnitude values were calculated:

Magnitude scale	Value
Mw (USGS)	6.2
mb (USGS)	5.7
mb	6.4
mB_BB	7.0
Ms_20	6.3
Ms_BB	6.1

Note that Ms_20 and Ms_BB are within 0.1 m.u. of the USGS Mw. In contrast, both our mb and mB_BB are significantly larger, with mb, based on only 4 measurements in the near teleseismic distance range between 2000 and 3000 km, even by 0.7 m.u. larger than USGS mb, based on 299 measurements at a much larger range of distances up to about 11,000 km. This may hint to larger calibration uncertainties of the Gutenberg-Richter (1956) Q function in the near teleseismic distance range, where wave propagation is largely influenced by the greater inhomogeneity of upper mantle wave propagation and attenuation. This may result in larger uncertainty of body-wave magnitude estimates, if they are available in only very limited number in the near teleseismic range. On the other hand, this earthquake has taken place on the North-Atlantic Ridge (see Google map below). Intra-oceanic ridge strike-slip earthquakes are often high stress drop events characterized by a much larger than average ratio of E_S/M_0 and thus much larger M_e than Mw (e.g. Bormann and Di Giacomo, 2011; Choy, 2012, in IS 3.5). Then, mb and mB_BB may be much larger than usual too in their relation to long-period Mw and Ms. Therefore one might consider this as an alternative explanation, which is, however, neither supported by the much more representative USGS mb value (being rather normal for Mw = 6.2) nor by the relatively long and not very impulsive P waveforms. The latter are, in the case of high stress drop events, expected to be relatively short, impulse-like and of more high-frequency content (e.g., Fig. 3.6 in Chapter 3). Instead, the P-wave group is rather long for an Mw6.2 event, almost one minute, and relatively long-period.



Location map in GoogleEarth: This map shows the location with its error ellipse on the Mid-Atlantic ridge and the stations that were used for the analysis and location.

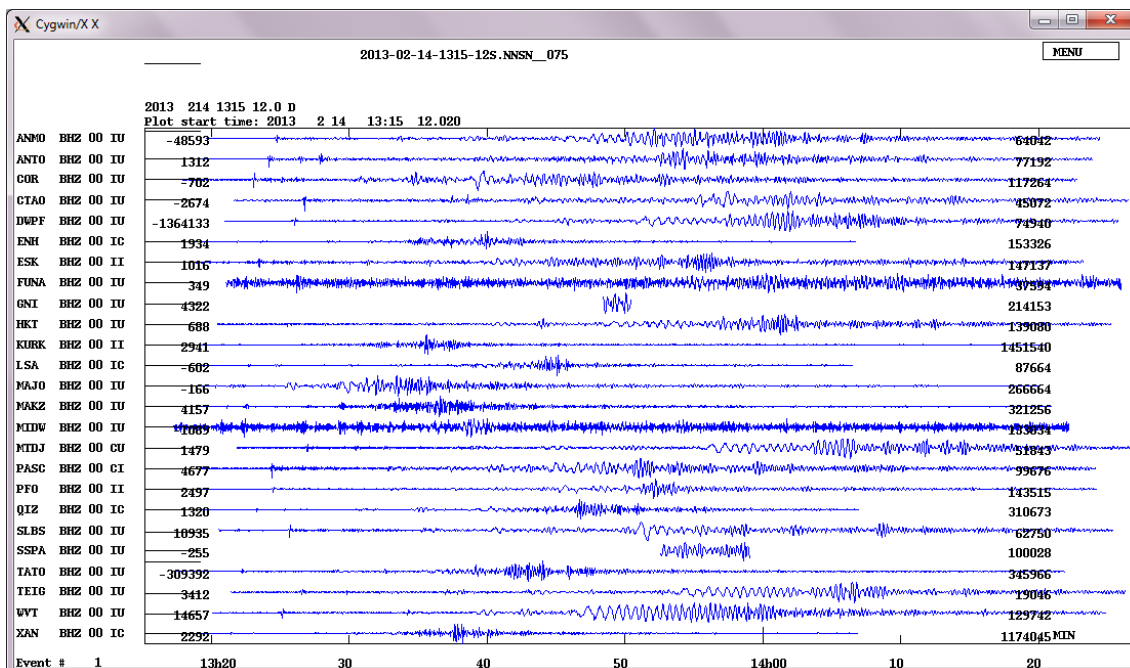
4 Analysis of a shallow earthquake recorded at teleseismic distances ($30^\circ < \Delta < 90^\circ$)

The earthquake chosen here as an example occurred in Russia at a shallow depth of 10 km (according to USGS). We extracted data from IRIS that was recorded on the Global Seismograph Network (GSN). All data used here were recorded on broadband stations.

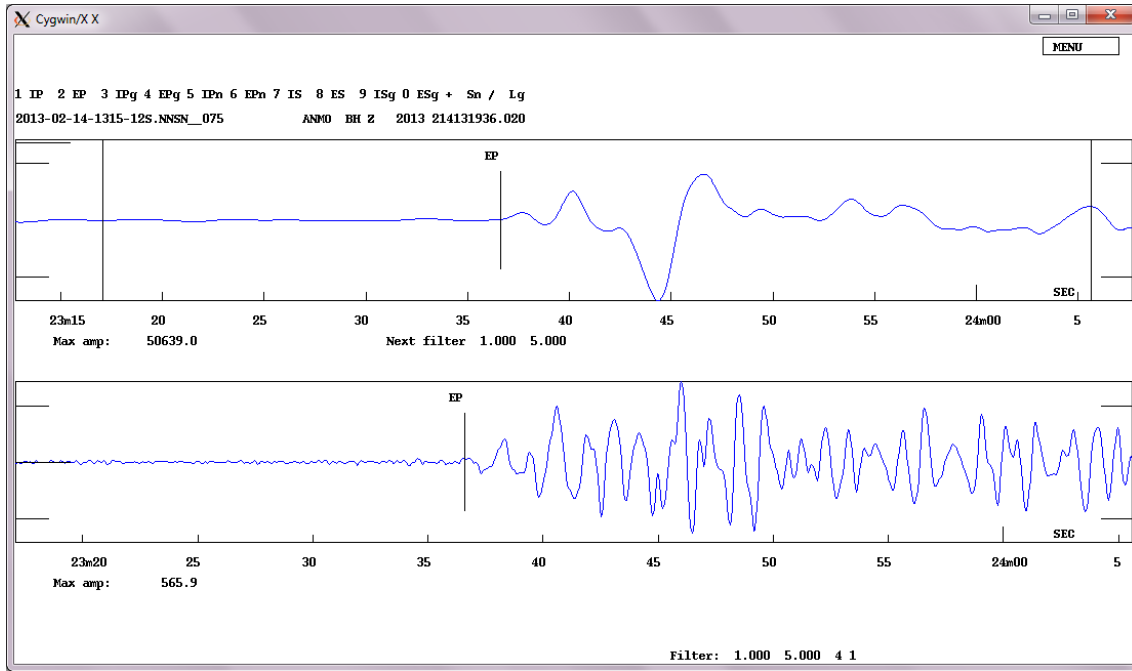
The earthquake parameters provided by the USGS are:

Date:	14/02/2013
Origin time:	13:13:53.1 (UTC)
Latitude:	67.61°N
Longitude:	142.60°W
Depth:	10 km
Mw:	6.9

We now show the step-by-step procedure to analyze the earthquake.



Multi-trace plot of vertical components, unfiltered: Record durations up to about 60 minutes at some stations and the observation of much larger surface-wave than body wave amplitudes hints to a large shallow earthquake at teleseismic distances.



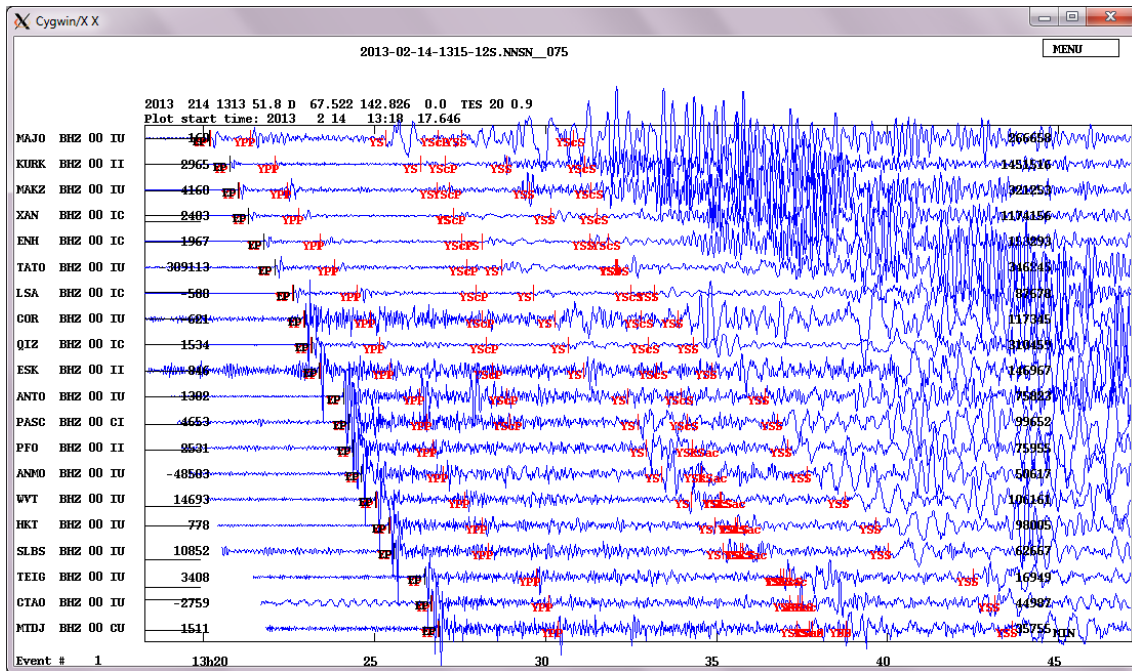
Vertical-component trace of station ANMO at $\Delta = 65.4^\circ$, unfiltered (top) and filtered 1-5 Hz (bottom): We start by reading the P arrivals that are very clear for this event. They can be read both on unfiltered (top) and high-pass filtered seismograms (bottom).

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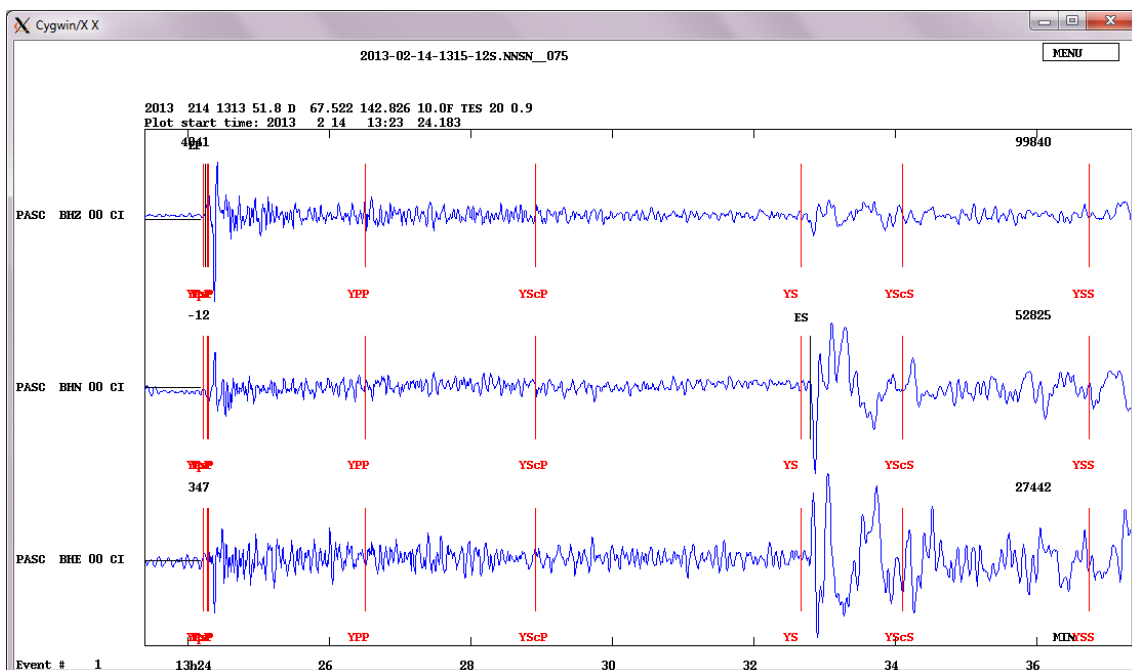
# 1 14 Feb 2013 13:15 12 D ? 1
date hrmn sec lat long depth no m rms damp erln erlt erdp
13 214 1313 51.82 6731.31N 142 49.6E 0.0 19 3 0.94 0.000 42.6 19.3236.1
stn dist azm ain w phas calcpchs hrmn tsec t-obs t-cal res wt di
MAJO 3456 187.2 27.3 0 P P 1320 10.0 378.22 379.99 -1.77 0.90*12
KURK 3888 274.1 26.7 0 P P 1320 45.3 413.45 413.76 -0.31 1.00* 9
MAKZ 4098 267.2 26.3 0 P P 1321 1.0 429.15 430.05 -0.90 0.97* 7
XAN 4307 227.7 25.9 0 P P 1321 18.1 446.28 446.03 0.25 1.00* 5
ENH 4683 225.1 25.2 0 P P 1321 45.2 473.41 474.06 -0.65 0.99* 5
TATO 4945 208.1 24.6 0 P P 1322 5.6 493.80 493.00 0.80 0.98* 7
LSA 5384 245.7 23.7 0 P P 1322 38.1 526.24 524.80 1.44 0.93* 3
COR 5678 66.3 23.1 0 P P 1322 57.4 545.60 544.19 1.41 0.94* 4
QIZ 5863 220.3 22.7 0 P P 1323 9.3 557.43 556.76 0.67 0.99* 5
ESK 6093 337.0 22.2 0 P P 1323 23.0 571.15 571.46 -0.31 1.00*11
ANTO 6748 304.1 20.8 0 P P 1324 5.6 613.77 613.41 0.36 1.00* 9
PASC 6905 67.8 20.5 0 P P 1324 14.8 622.99 623.05 -0.06 1.00* 4
PFO 7029 66.6 20.3 0 P P 1324 22.4 630.63 630.69 -0.06 1.00* 4
ANMO 7271 57.6 19.7 0 P P 1324 36.6 644.80 644.94 -0.14 1.00* 3
WVT 7753 41.8 18.7 0 P P 1325 2.0 670.19 672.03 -1.84 0.89* 2
HKT 8152 50.7 17.9 0 P P 1325 25.0 693.21 693.77 -0.56 0.99* 3
SLBS 8298 65.2 17.6 0 P P 1325 33.8 701.95 701.74 0.21 1.00* 3
TEIG 9414 47.2 15.2 0 P P 1326 27.4 755.63 756.06 -0.43 0.99* 2
CTAO 9700 176.8 14.5 0 P P 1326 39.3 767.43 768.26 -0.83 0.00 0
MTDJ 9936 38.0 14.1 0 P P 1326 52.7 780.87 778.91 1.96 0.88* 1

2013 214 1313 51.8 D 67.522 142.826 0.0 TES 20 0.9
    
```

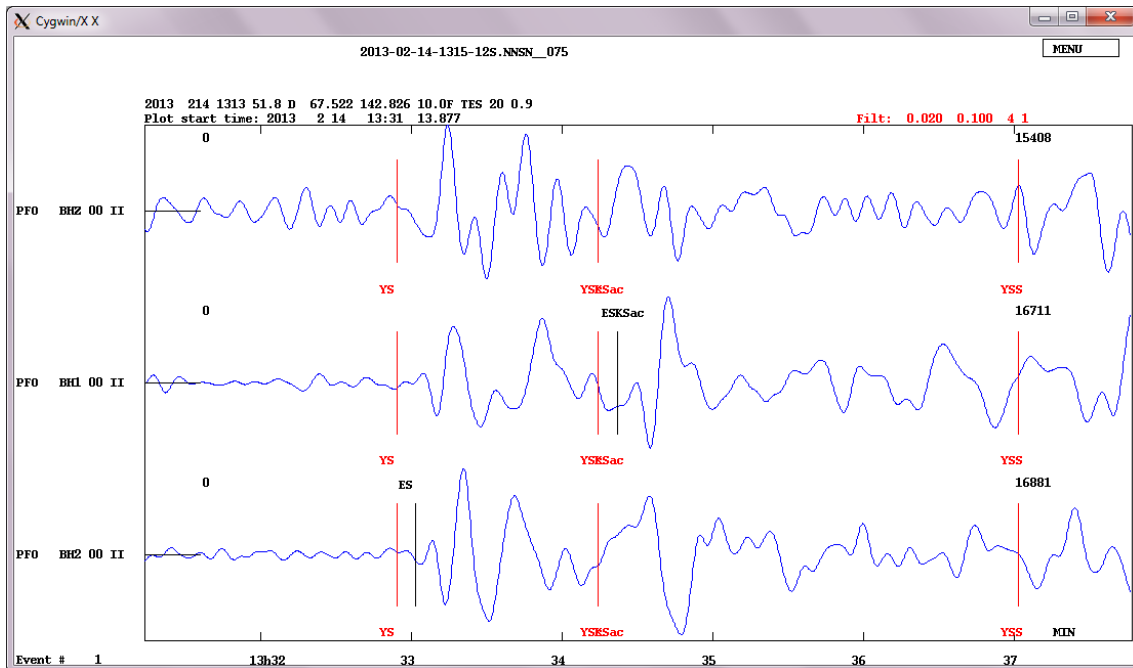
Output from the location program: Locating with the first P arrivals only, we obtain an epicenter location (67.522°N , 142.826°W) and depth = 0 km that is very close to the one given by the USGS. The depth is not constrained with the teleseismic P arrivals only.



Multi-trace plot of vertical components, unfiltered: Plotted are the traces on which we have read the initial P. Based on the hypocenter obtained, we compute the synthetic arrival times of all phases. The traces are sorted by distance. The phases PP, S and SS are clear on all traces. We will now read the onset times of all these phases (plus P-wave first motion polarity), run the location program again with this additional information and see how the location changes.



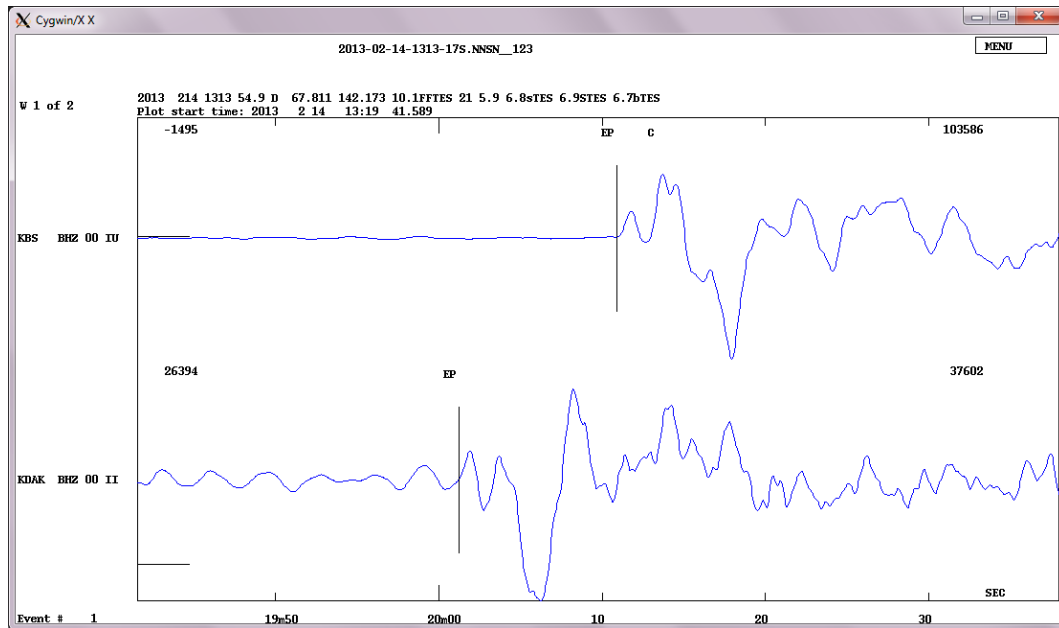
Three-component trace plot of station PASC at $\Delta=62.2^\circ$, unfiltered: For this station P on the vertical component and S on the N component are very clear. Other phases are not very clear. Note that we do not pick the onsets at the calculated arrival times (red bars) but only at the distinct onset of the real phases with large SNR (black bars).



Three-component trace plot of station PFO at $\Delta = 63.3^\circ$, filtered 0.02-0.1 Hz: At this distance there is sufficient separation between S and SKSac to read both phases. The component codes of the horizontal components here are BH1 and BH2 to indicate that the instrument is not exactly aligned with the geographic directions. In this case, the direction of orientation is given with the component's metadata and can be used to rotate into a NS-EW system.

#	1	14	Feb	2013	13:13	51	D	67.522	142.826	0.9	TES	20	? 1	
date	hrmn	sec	lat	long	depth	no	m	rms	damp	erln	erlt	erdp		
13 214	1313	54.89	6748.67N 142	10.4E	0.1	55	3	5.94	0.000	92.7	44.2	39.3		
stn	dist	azm	ain	w	phas	calcp	phs	hrmn	tsec	t-obs	t-cal	res	wt	di
MAJO	3485	186.1	42.8	0	PP	PnPn		1321	22.8	447.91	445.00	2.91	1.00	4
MAJO	3485	186.1	27.3	0	P	P		1320	10.0	375.15	382.25	-7.11	1.00	2
MAJO	3485	186.1	28.1	0	S	S		1325	19.8	684.86	691.48	-6.62	1.00	6
KURK	3858	273.1	26.8	0	P	P		1320	45.3	410.38	411.44	-1.06	1.00	1
KURK	3858	273.1	40.6	0	PP	PnPn		1322	2.6	487.69	487.73	-0.04	1.00	3
KURK	3858	273.1	27.7	0	S	S		1326	18.9	744.06	743.37	0.69	1.00	4
MAKZ	4072	266.2	26.4	0	P	P		1321	1.0	426.08	428.05	-1.98	1.00	1
MAKZ	4072	266.2	39.7	0	PP	PnPn		1322	27.4	512.54	511.60	0.94	1.00	2
MAKZ	4072	266.2	46.9	0	SS	SnSn		1329	16.4	921.49	924.72	-3.23	1.00	9
XAN	4308	226.6	25.9	0	P	P		1321	18.1	443.21	446.13	-2.92	1.00	0
XAN	4308	226.6	34.8	0	PP	PP		1322	52.7	537.78	535.32	2.46	1.00	1
XAN	4308	226.6	27.1	0	S	S		1327	16.3	801.38	805.45	-4.08	1.00	2
ENH	4687	224.0	25.2	0	P	P		1321	45.2	470.34	474.30	-3.96	1.00	0
ENH	4687	224.0	34.3	0	PP	PP		1323	30.7	575.79	572.36	3.43	1.00	1
ENH	4687	224.0	26.5	0	S	S		1328	6.2	851.27	856.26	-4.99	1.00	2
ENH	4687	224.0	36.7	0	SS	SS		1331	20.0	1045.07	1046.08	-1.01	1.00	3
TATO	4960	207.1	24.6	0	P	P		1322	5.6	490.73	494.12	-3.39	1.00	1
TATO	4960	207.1	33.8	0	PP	PP		1323	57.4	602.48	598.76	3.72	1.00	2
TATO	4960	207.1	26.0	0	S	S		1328	43.4	888.52	892.20	-3.68	1.00	3
LSA	5373	244.7	23.7	0	P	P		1322	38.1	523.17	523.96	-0.80	1.00	0
LSA	5373	244.7	28.4	0	PP	PP		1324	35.9	641.06	636.87	4.19	1.00	1
LSA	5373	244.7	25.3	0	S	S		1329	40.4	945.49	946.59	-1.10	1.00	1
LSA	5373	244.7	29.2	0	SS	SS		1333	24.0	1169.11	1158.32	10.78	1.00	2
COR	5690	66.0	23.1	0	P	P		1322	57.4	542.53	545.00	-2.47	1.00	1
COR	5690	66.0	24.7	0	S	S		1330	21.9	986.97	985.30	1.67	1.00	2
COR	5690	66.0	28.6	0	SS	SS		1334	25.7	1230.76	1203.01	27.74	1.00	3
QIZ	5870	219.4	22.7	0	P	P		1323	9.3	554.36	557.21	-2.85	1.00	0
QIZ	5870	219.4	28.1	0	PP	PP		1325	15.6	680.66	677.02	3.64	1.00	1
QIZ	5870	219.4	24.4	0	S	S		1330	41.4	1006.49	1007.80	-1.32	1.00	2
QIZ	5870	219.4	28.5	0	SS	SS		1334	45.3	1250.38	1228.62	21.76	1.00	2
ESK	6053	336.5	22.3	0	P	P		1323	23.0	568.08	568.79	-0.71	1.00	2
ANTO	6707	303.4	20.9	0	P	P		1324	5.6	610.70	610.87	-0.17	1.00	2
ANTO	6707	303.4	27.5	0	PP	PP		1326	22.1	747.22	744.41	2.81	1.00	3
ANTO	6707	303.4	0					1327	47.7	832.8				
ANTO	6707	303.4	22.8	0	S	S		1332	27.0	1112.14	1107.57	4.57	1.00	6
PASC	6918	67.4	20.5	0	P	P		1324	14.8	619.92	623.83	-3.91	1.00	1
PASC	6918	67.4	22.4	0	S	S		1332	47.8	1132.86	1131.89	0.97	1.00	2
PFO	7042	66.2	20.2	0	P	P		1324	22.4	627.56	631.41	-3.85	1.00	1
PFO	7042	66.2	22.2	0	S	S		1333	1.5	1146.59	1146.11	0.48	1.00	2
PFO	7042	66.2	13.3	0	SKSac	SKSac		1334	22.1	1227.24	1226.38	0.86	1.00	1
ANMO	7276	57.2	19.7	0	P	P		1324	36.6	641.73	645.27	-3.54	1.00	1
ANMO	7276	57.2	21.7	0	S	S		1333	24.0	1169.09	1172.25	-3.16	1.00	2
WVT	7747	41.4	18.8	0	P	P		1325	2.0	667.12	671.69	-4.57	1.00	1
WVT	7747	41.4	20.8	0	S	S		1334	11.7	1216.83	1222.45	-5.62	1.00	2
WVT	7747	41.4	27.7	0	SS	SS		1338	42.6	1487.75	1492.76	-5.01	1.00	3
HKT	8153	50.2	17.9	0	P	P		1325	25.0	690.14	693.79	-3.65	1.00	1
HKT	8153	50.2	20.0	0	S	S		1334	55.2	1260.27	1264.73	-4.46	1.00	2
SLBS	8309	64.7	17.6	0	P	P		1325	33.8	698.88	702.32	-3.44	1.00	1
SLBS	8309	64.7	19.7	0	S	S		1335	15.7	1280.78	1281.06	-0.29	1.00	2
SLBS	8309	64.7	27.3	0	SS	SS		1340	12.2	1577.30	1570.17	7.13	1.00	3
TEIG	9412	46.6	15.2	0	P	P		1326	27.4	752.56	755.95	-3.39	1.00	0
TEIG	9412	46.6	25.1	0	PP	PP		1329	48.6	953.66	952.02	1.64	1.00	1
CTAO	9733	176.2	14.4	0	P	P		1326	39.3	764.36	769.71	-5.35	0.00	0
CTAO	9733	176.2	24.8	0	PP	PP		1330	6.6	971.71	975.50	-3.79	1.00	2
MTDJ	9927	37.4	10.3	0	SKSac	SKSac		1337	24.1	1409.23	1409.91	-0.68	1.00	1
MTDJ	9927	37.4	24.6	0	PP	PP		1330	25.7	990.79	989.60	1.18	1.00	1
MTDJ	9927	37.4	14.1	0	P	P		1326	52.7	777.80	778.52	-0.72	1.00	0
2013	214	1313	54.9	D	67.811	142.173	0.1	TES	20	5.9				

Output from the location program: Using additional phases the location changes slightly, but without readings of depth phases or P-wave onsets very near to this shallow source the depth cannot be better resolved compared to the USGS.

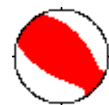


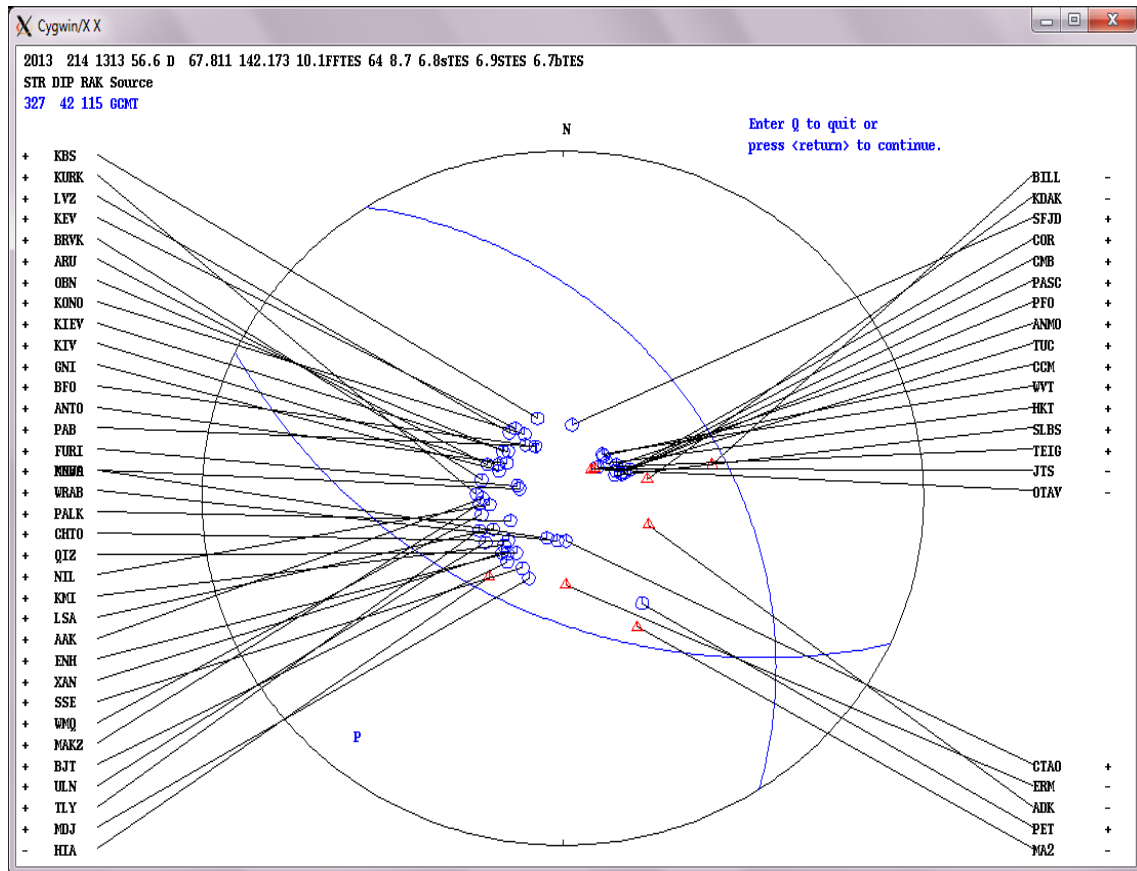
Vertical component traces of stations KBS and KDAK, unfiltered: We try to read first-motion P-wave polarities (upward motion = C = + = compressional, downward motion = D = - = dilatational) which can allow us to estimate a fault-plane solution for this earthquake. For station KBS (top) it is possible to read the polarity, which is upward (compression). For station KDAK it is not possible to read the polarity, although we were tempted to read this as a downward motion (dilatation). As seen further down, the selected data set does not have stations that are for sure in a dilatational quadrant, however, the position of all our compressional readings (see next figure) are in agreement with Global CMT fault-plane solution.

The following solution was provided by the Global CMT catalog (www.globalcmt.org) for this event. The solution shows an oblique thrust mechanism:

201302141313A EASTERN SIBERIA, RUSSIA

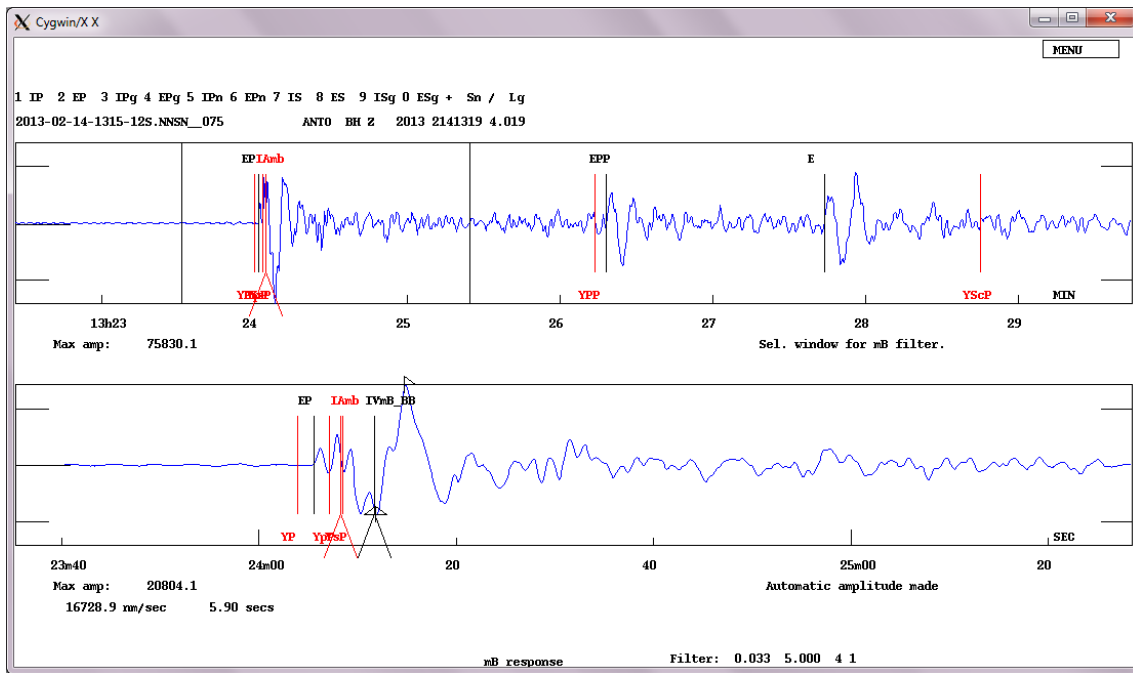
Date: 2013/ 2/14 Centroid Time: 13:13:58.6 GMT
 Lat= [67.71](#) Lon= [142.62](#)
 Depth= 12.0 Half duration= 5.2
 Centroid time minus hypocenter time: 5.5
 Moment Tensor: Expo=26 1.030 -0.627 -0.404 0.335 0.075 0.729
 Mw = 6.7 mb = 0.0 Ms = 6.6 Scalar Moment = 1.2e+26
 Fault plane: strike=327 dip=42 slip=115
 Fault plane: strike=115 dip=52 slip=69



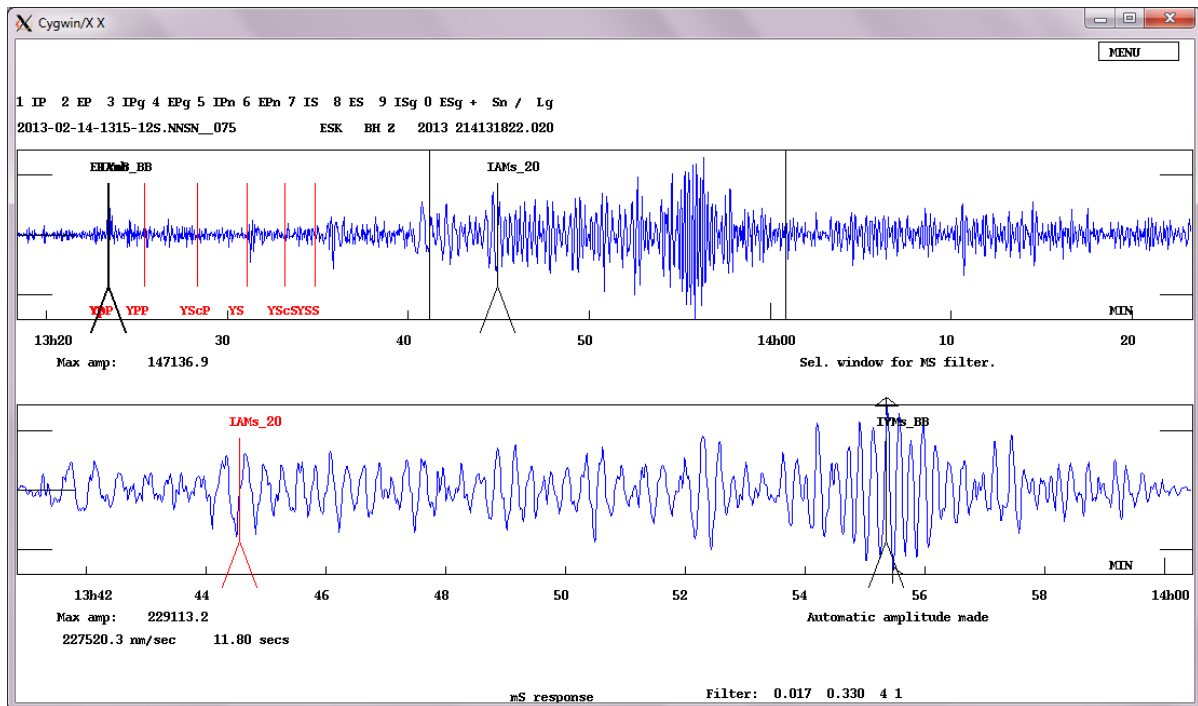


Fault-plane solution plot: This plot is created in SEISAN and shows the solution given by the Global CMT catalog together with polarities read from the example. We can see that all our observations lie near the center of the focal sphere and are likely compressions. However, the readings we made show compression and dilatation, and most likely the dilatational readings are wrong. Although looking at more than 100 stations from the global networks, we do not have good enough coverage of the focal sphere as all our rays to teleseismic stations only leave the source downward near the vertical. In contrast, the moment tensor solution (based on full waveforms from P, S and surface waves) normally requires only few stations to obtain a good solution. Using complementary to first-motion polarity readings in SEISAN also S/P amplitude ratios for better constraining fault-plane solutions is possible for local and teleseismic earthquakes. While the use of amplitude ratios did not help with this example, Figure 7.26 in Havskov and Ottemöller (2010) gives an example of teleseismic fault plane solution using amplitude ratios. The polarity readings made by us on individual station records (open circles = C, triangles = D) are given on the side with a link to the azimuth and angle of incidence of the corresponding phase projected onto the focal sphere.

Finally, we demonstrate the magnitude determinations for this earthquake:



Traces of station ANTO at $\Delta = 60.3^\circ$, *mB_BB* filter: First we read amplitudes for mb and mB_BB. Measured are the largest P-wave amplitude prior to PP, depth phases of P included, on the vertical component of simulated short-period WWSSN, respectively velocity broadband records. The position of the earlier Iamb measurement has been plotted here on the mB_BB filtered trace only for comparison by a red arrow-bar. Details of teleseismic body-wave magnitudes are given in the regional event example in section 3 above and in IS 3.3.



Traces of station ESK at $\Delta = 54.4^\circ$, *Ms_BB* filter: Next we determine the two IASPEI standard surface-wave magnitudes. The bottom trace shows the reading position of the *Ms_BB* amplitude *IVMs_BB* (black arrow-bar) on an instrument corrected broadband velocity trace.

The top trace is unfiltered and shows the time selection for the bottom trace. Note that the IAMS_20 amplitude has already been read on another, WWSSN-LP filtered, record, before. However, the position of its measurement is shown here too for comparison (marked by a red arrow-bar). While the largest Ms_20 displacement amplitude has a period of 21 s, the period of the 10.5 min later arriving largest Ms_BB velocity amplitude is 11.80 s only. For more details on the determination of teleseismic surface-wave magnitude see the regional event example in section 3 as well as IS 3.3.

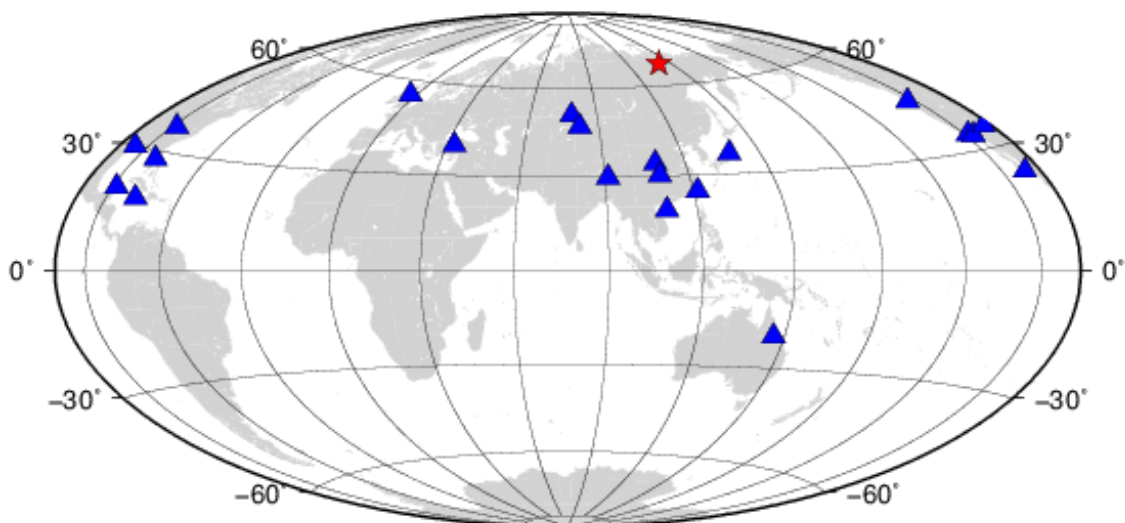
Computing the average magnitudes from all amplitude observations we get the values below:

Magnitude scale	Value
Mw (USGS)	6.9
mb	6.7
mB_BB	7.3
Ms_20	6.9
Ms_BB	6.9

In this case the surface wave magnitudes Ms_20 and Ms_BB agree perfectly with Mw(USGS) and are equal amongst each other although Ms_BB has been measured at a much shorter period and a later time than Ms_20. mB_BB is significantly larger than mb, and somewhat larger than the two Ms and the Mw magnitude, as expected for earthquakes of Mw = 6.9 according to Utsu (2002) and Fig. 3.70 in section 3.2.9.1 of the forthcoming revised Chapter 3).

The amplitudes and periods that were read for computing the magnitudes are summarized in the following table. Note that here the abridged internal magnitude nomenclature used in SEISAN stands for: mb = mb, mB = mB_BB, Ms = Ms_20 and MS = Ms_BB.

XAN	BZ	dist:	4308.0	amp:	1498.7	T:	1.8	mb =	6.3
XAN	BZ	dist:	4308.0	amp:	22757.5	T:	8.8	mB =	7.0
ENH	BZ	dist:	4687.0	amp:	959.5	T:	1.5	mb =	6.3
ENH	BZ	dist:	4687.0	amp:	17567.1	T:	8.8	mB =	6.9
ENH	BZ	dist:	4687.0	amp:	120000.0	T:	10.9	MS =	7.3
ENH	BZ	dist:	4687.0	amp:	184000.0	T:	18.6	Ms =	7.0
TATO	BZ	dist:	4960.0	amp:	436.4	T:	0.9	mb =	6.3
TATO	BZ	dist:	4960.0	amp:	20681.6	T:	9.0	mB =	7.2
TATO	BZ	dist:	4960.0	amp:	111000.0	T:	18.9	MS =	6.8
TATO	BZ	dist:	4960.0	amp:	81389.9	T:	11.8	MS =	7.2
COR	BZ	dist:	5690.0	amp:	18587.7	T:	3.8	mB =	7.2
COR	BZ	dist:	5690.0	amp:	778.2	T:	1.1	mb =	6.5
COR	BZ	dist:	5690.0	amp:	11899.5	T:	15.7	MS =	6.4
COR	BZ	dist:	5690.0	amp:	31700.7	T:	18.9	MS =	6.4
ESK	BZ	dist:	6053.0	amp:	6275.4	T:	1.1	mb =	7.6
ESK	BZ	dist:	6053.0	amp:	171000.0	T:	8.5	mB =	8.2
ESK	BZ	dist:	6053.0	amp:	336000.0	T:	20.8	MS =	7.4
ESK	BZ	dist:	6053.0	amp:	228000.0	T:	11.8	MS =	7.7
ANTO	BZ	dist:	6707.0	amp:	721.5	T:	1.2	mb =	6.7
ANTO	BZ	dist:	6707.0	amp:	16728.9	T:	5.9	mB =	7.3
ANTO	BZ	dist:	6707.0	amp:	57305.5	T:	19.2	MS =	6.7
ANTO	BZ	dist:	6707.0	amp:	20010.4	T:	20.0	MS =	6.8
ANMO	BZ	dist:	7276.0	amp:	9820.2	T:	3.8	mB =	7.2
ANMO	BZ	dist:	7276.0	amp:	892.4	T:	2.0	mb =	6.6
ANMO	BZ	dist:	7276.0	amp:	18195.2	T:	21.0	MS =	6.8
ANMO	BZ	dist:	7276.0	amp:	59228.9	T:	21.3	MS =	6.8
WVT	BZ	dist:	7747.0	amp:	68435.1	T:	20.2	MS =	6.9
WVT	BZ	dist:	7747.0	amp:	20200.8	T:	19.0	MS =	6.9
DWPF	BZ	dist:	8770.0	amp:	1956.5	T:	2.2	mb =	6.8
DWPF	BZ	dist:	8770.0	amp:	11188.1	T:	3.6	mB =	7.1
DWPF	BZ	dist:	8770.0	amp:	74362.7	T:	22.0	MS =	7.0
DWPF	BZ	dist:	8770.0	amp:	22267.2	T:	23.6	MS =	7.0
TEIG	BZ	dist:	9412.0	amp:	12224.9	T:	2.9	mB =	7.3
TEIG	BZ	dist:	9412.0	amp:	853.1	T:	1.4	mb =	6.8
TEIG	BZ	dist:	9412.0	amp:	18134.1	T:	21.4	MS =	7.0
TEIG	BZ	dist:	9412.0	amp:	58385.9	T:	20.6	MS =	7.0
CTAO	BZ	dist:	9733.0	amp:	686.2	T:	2.0	mb =	6.6
CTAO	BZ	dist:	9733.0	amp:	4437.9	T:	19.5	MS =	6.4
CTAO	BZ	dist:	9733.0	amp:	13972.4	T:	21.4	MS =	6.3



Map showing the epicenter (red star) and stations (blue triangles) used here.

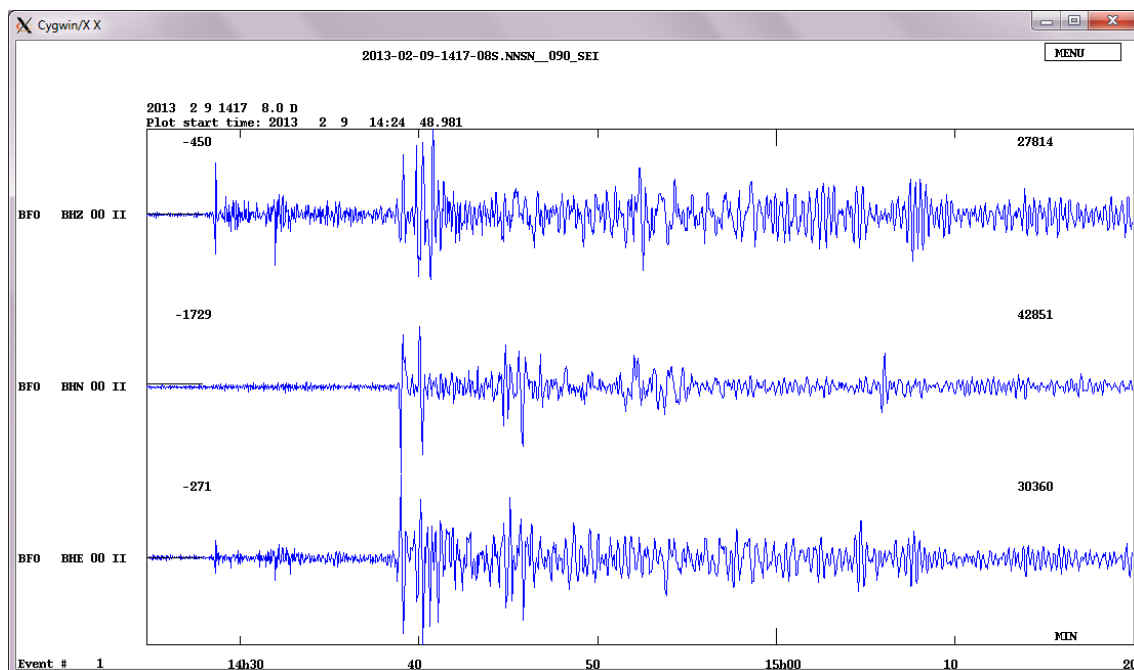
5 Analysis of an intermediate depth earthquake at teleseismic distances

The earthquake chosen here as example occurred in Colombia at an intermediate depth of 129 km (depth by USGS). We extracted data from IRIS that were recorded by the Global Seismograph Network (GSN). All data used here were recorded on broadband stations.

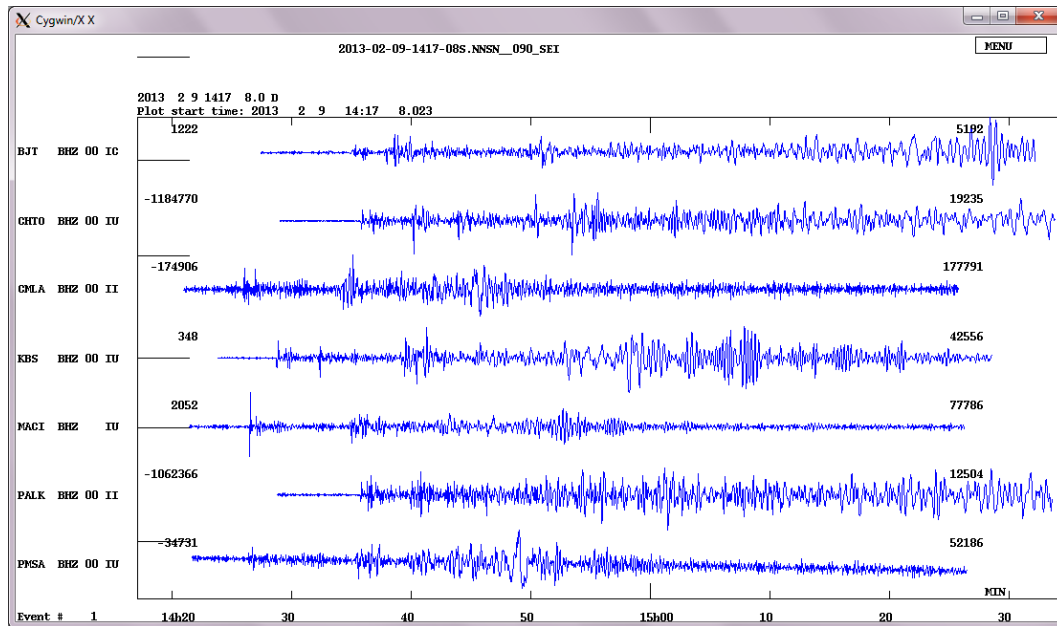
The earthquake parameters provided by the USGS are:

Date:	09/02/2013
Origin time:	14:16:06.3 (UTC)
Latitude:	1.14°N
Longitude:	77.36°W
Depth:	129 km
Mw:	7.0

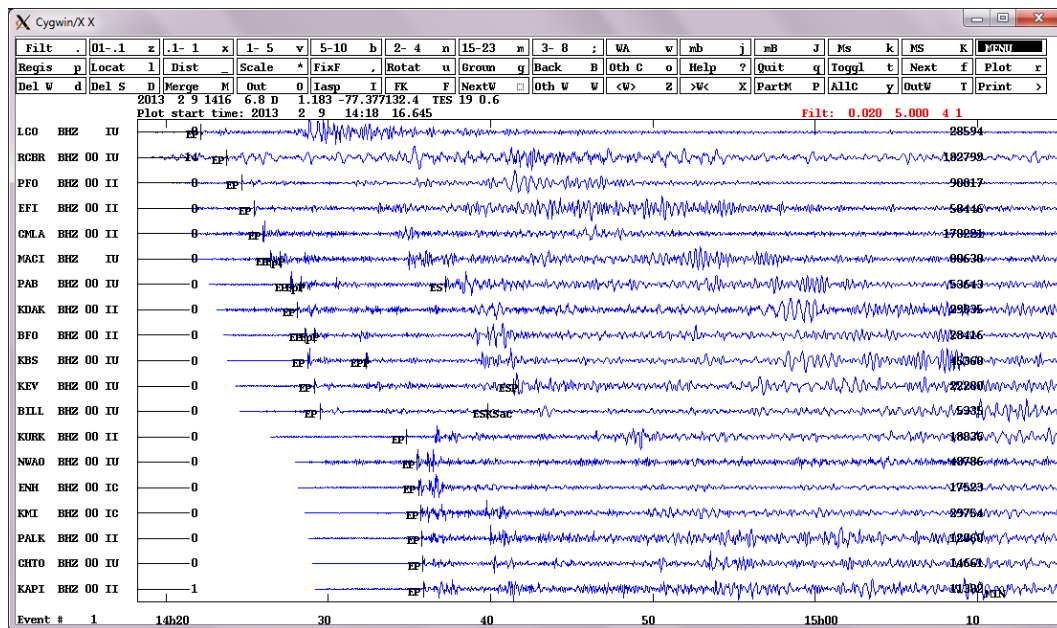
We now show the step-by-step procedure to analyze the earthquake.



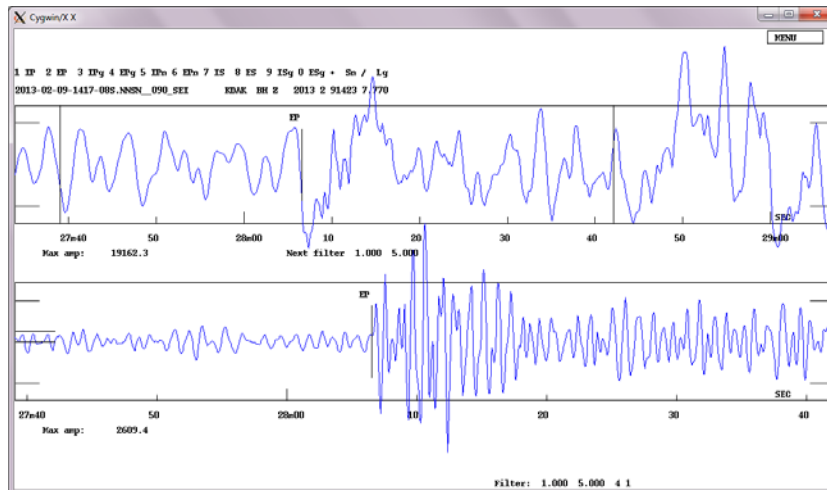
Three-component trace plot of station BFO at $\Delta = 86.4^\circ$, unfiltered: Looking at one of the stations, we can see that we are dealing with an earthquake at teleseismic distance as the record shown extends to almost one hour. The seismograms show very clear body-wave phase arrivals, which we will try to identify. The lack of more long-period and dispersed surface waves, however, points at a deep hypocenter. We will start to pick first P arrivals to get an initial estimate of the location, then we try to identify depth phases and other later phases for improved hypocenter determination and finally we use the 3-component record of station BFO only for a single-station event location.



Multi-trace plot of vertical components, unfiltered: Shown is only a selection of the stations that are part of the sample data set for this event.



Vertical-component traces for which phases are picked, filtered 0.02-5 Hz: These traces are sorted by epicentral distance, ranging from 30.7° to 162.7° . Beside the P first arrival, at some of the traces also clear later arrivals have been marked and identified. Note that from station KURK downward the strong delay of the first arriving longitudinal waves is due to the Earth core shadow for direct P waves. In fact, the group of later arrivals, annotated here as P, relate to different branches of PKP core phases and likely also their depth phases. Also note several distinct onsets following within the first minute after the P onset and well before the arrival of PP. For a first event location we will only use the P and PKP first arrivals. Later, we will zoom into the first few minutes of records in order to find depth phases of P and PKP.

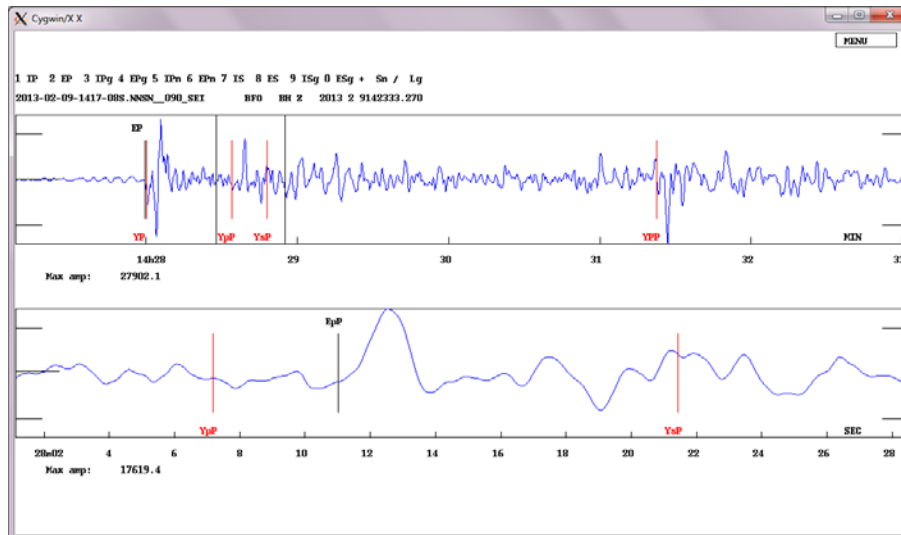


Vertical-component trace of station KDAK at $\Delta=81.2^\circ$, top unfiltered, bottom filtered 1-5 Hz: Whereas most of the P arrivals in the earlier shown record section of all stations had a good SNR, this is an example where the P arrival can only be picked after filtering the trace. The time window of the expanded lower trace is marked by the vertical lines on the unfiltered top trace. The use of the filter introduces a small phase shift and at some observatories it is common practice to read all stations with the same filter. However, as long as the phase shift is small as compared to the reading error one can, depending on the signal-to-noise ratio (SNR) of individual stations, chose the best filter or no filter. There may also be stations where the SNR is too low and phases will not be read. This is better than reading false phases.

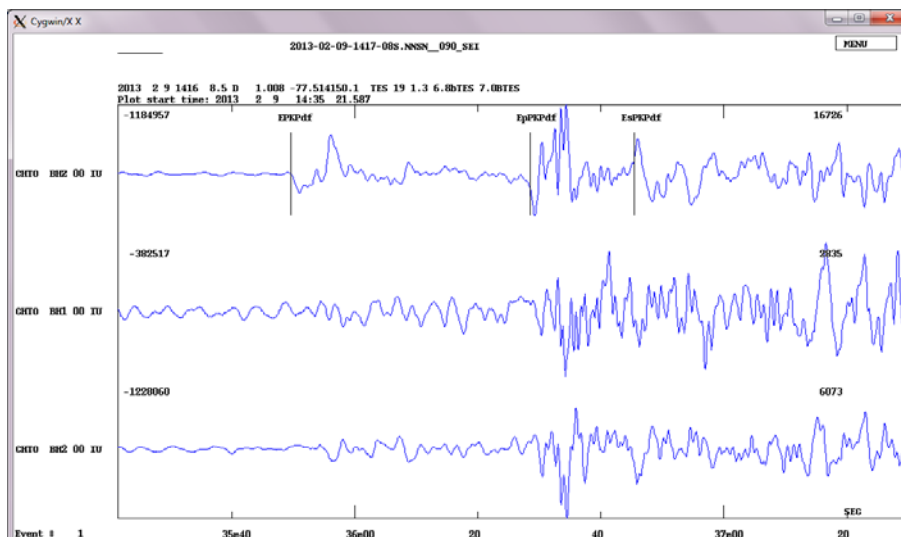
```
# 1 9 Feb 2013 14:17 8 D ? 1
date hrmn sec lat long depth no m rms damp erln erlt erdp
13 2 9 1416 6.75 110.98N 77 22.6W 132.4 19 3 0.64 0.000 15.6 12.3 48.7
stn dist azm ain w phas calcphs hrmn tsec t-obs t-cal res wt di
LCO 3417 168.5 39.7 0 P P 1422 9.6 362.83 362.82 0.01 1.00*17
RCBR 4675 100.0 36.2 0 P P 1423 45.0 458.23 458.28 -0.05 1.00*11
PFO 5430 315.6 33.7 0 P P 1424 39.4 512.70 511.80 0.90 0.97*21
EFI 6139 165.5 31.5 0 P P 1425 25.1 558.33 558.79 -0.47 0.99* 9
CMLA 6665 46.1 29.8 0 P P 1425 59.0 592.23 591.93 0.30 1.00* 6
MACI 7120 59.1 28.4 0 P P 1426 27.2 620.43 619.48 0.95 0.97* 5
PAB 8482 49.6 24.3 0 P P 1427 40.6 693.81 694.00 -0.20 1.00* 3
KDAK 9032 328.4 22.6 0 P P 1428 6.6 719.81 720.90 -1.09 0.96* 5
BFO 9603 41.8 20.8 0 P P 1428 32.9 746.20 747.04 -0.84 0.98* 2
KBS 9876 11.1 19.9 0 P P 1428 45.8 759.07 758.66 0.41 0.99* 1
KEV 10448 19.8 19.4 0 P P 1429 8.0 781.28 782.45 -1.18 0.96* 1
BILL 10966 340.1 18.8 0 P Pdif 1429 29.2 802.41 803.17 -0.76 0.98* 2
KURK 13847 18.3 8.0 0 P PKPdf 1434 49.31122.541123.03 -0.49 0.99* 2
NWA0 16204 202.0 7.0 0 P PKPdf 1435 28.61161.821161.84 -0.02 1.00* 4
ENH 16480 348.7 6.8 0 P PKPdf 1435 33.21166.431165.98 0.45 0.99* 2
KMI 17127 359.8 5.9 0 P PKPdf 1435 42.81176.031175.09 0.94 0.97* 2
PALK 17430 68.7 5.4 0 P PKPdf 1435 45.81179.101178.53 0.57 0.99* 3
CHTO 17789 10.1 4.8 0 P PKPdf 1435 49.51182.781182.46 0.32 1.00* 2
KAPI 18088 257.1 4.2 0 P PKPdf 1435 52.21185.461185.29 0.16 1.00* 3

2013 2 9 1416 6.8 D 1.183 -77.377132.4 TES 19 0.6
```

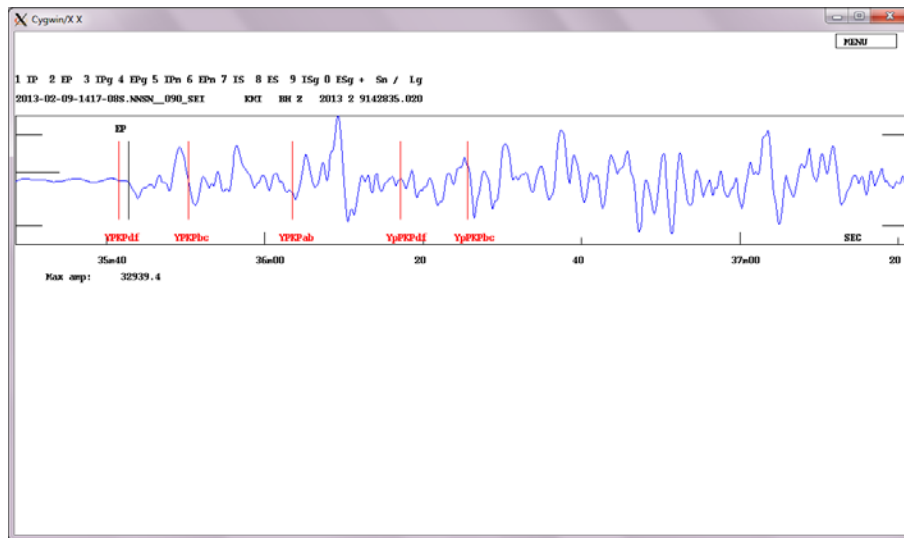
Output from the location program on the previous page: Locating with the first P arrivals only, we obtain a hypocenter location (1.183°N , 77.377°W and depth = 132.4 km) that is very close to the one given by the USGS. The depth estimate is perhaps better than expected in this case without the use of depth phases. However, we will now try to use the depth phases we noticed on a number of seismograms. We have so far read all first arrivals as P. The location program determines which phase arrives first. In our example we have P (until station KEV), Pdif (station BILL) and PKPdf (from station KURK onward).



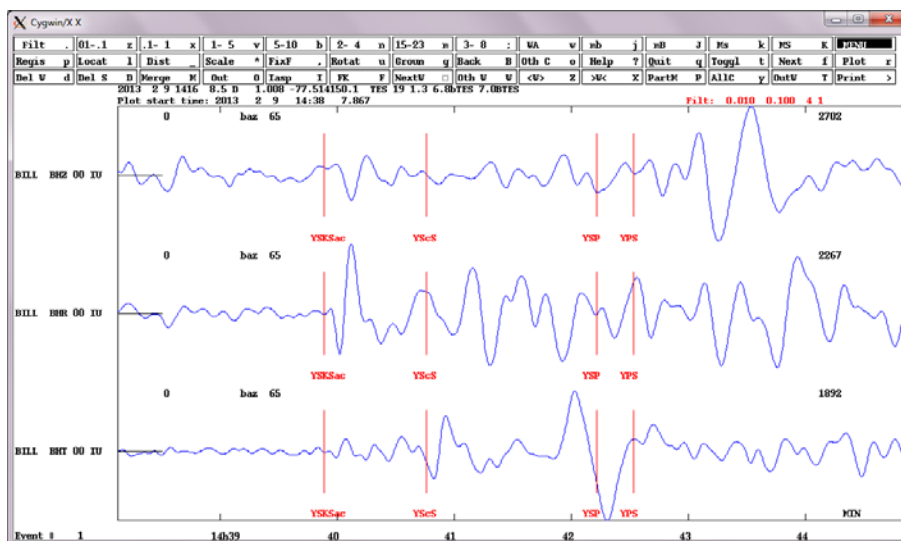
Vertical-component trace for station BFO at $\Delta = 86.4^\circ$, unfiltered: For this station we have computed synthetic arrival times based on our initial location, which are marked with red bars. We can see distinct phases arriving near the computed pP, sP and PP times. The bottom trace shows the picking of pP about 4 s after the predicted arrival, but there is no clear new energy arrival around the expected sP onset. Note, however, that sP may sometimes be even a stronger phase than pP. The amplitude relationship between these two depth phases depends on the radiation pattern of the source mechanism, which is different for P and S waves (see Chapter 3, section 3.4.1, Figs. 3.100-104), and on the different take-off angles of pP and sP from the source (see Chapter 2, section 2.6.3 with figures). From the measured travel-time difference of 38.2 s between pP and P we estimate a source depth of about 150 km, some 20 km deeper than the source depth calculated by the USGS.



Three-component trace of station CHTO at $\Delta = 160.0^\circ$, unfiltered: Knowing from the first location of the earthquake that the station is 160° away the first arrival at 35m50s is clearly not P, as originally marked, but PKPdf. The larger onset about 5 s later is a second phase of the same event, yet the much stronger sharp wave onset at about 36m28s is obviously the depth phase pPKPdf and the onset at about 36m45s the depth phase sPKPdf. Note that the travel-time difference pPKPdf-PKPdf of 38 s perfectly agrees with the time difference pP-P = 38.2 s measured at station BFO. Note that the differential times such as pP-P and sP-P should be near constant over a distance range.



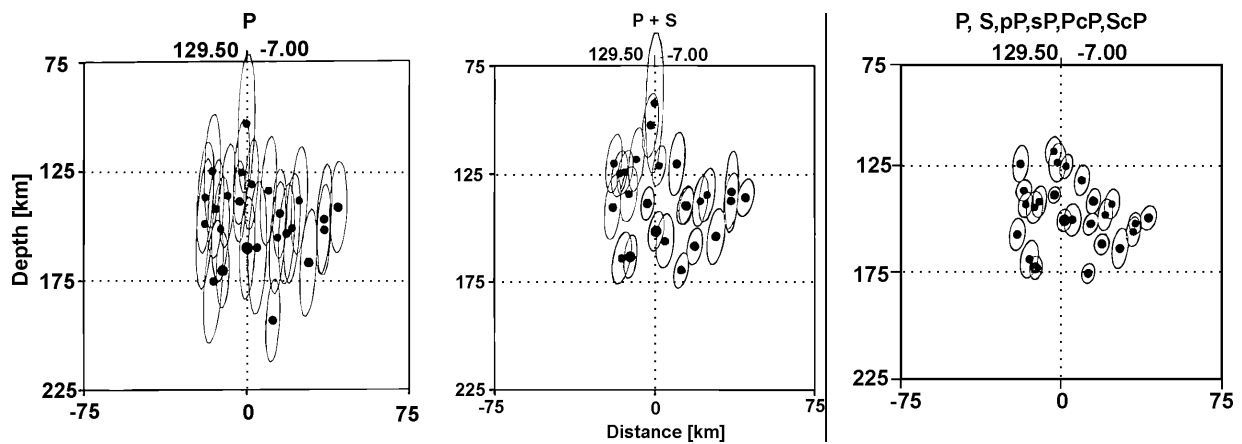
Vertical-component trace for station KMI at $\Delta=154.0^\circ$, unfiltered: In this example we can see onsets that relate to the different travel-time branches of PKP. The corresponding depth phase times are also shown and phases are visible. Core phases are often most clear in the frequency band 1-5 Hz.



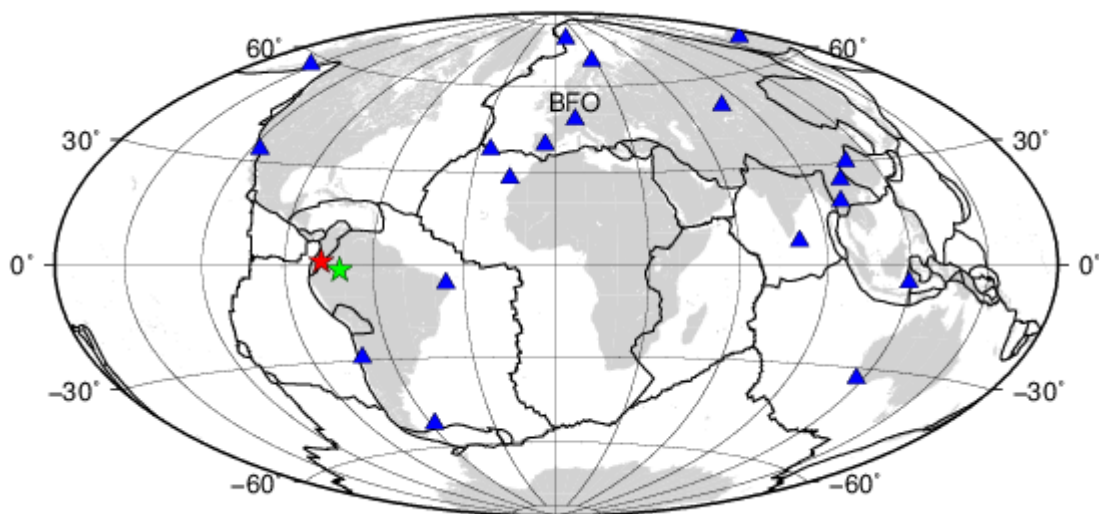
Three-component traces for station BILL at $\Delta = 98.6^\circ$, rotated and filtered 0.01-0.1 Hz: The two horizontal components have been rotated into radial and transverse components using a backazimuth of 65° . SKS has been back-converted from a P wave through the outer core, termed K, into an S wave through the mantle and is, therefore, polarised in the vertical plane of propagation. Accordingly, it has its largest amplitude in the radial component R (middle trace) and not on the transverse component T. It is also visible on the vertical component. This is not the rule for teleseismic S and illustrates that rotating seismograms can be useful in identifying seismic phases (see Figs. 11.12, 11.37 and 11.56 ?? in the Manual Chapter 11 of Bormann, 2009). Having an array of stations, the different apparent velocities of seismic phases can also be used for identification (see, e.g., Chapter 9, and for the identification of core phases Fig. 11.34 ?? in the Manual edition by Bormann (2009).

#	1	9 Feb 2013 14:16	8	D	1.008	-77.514150.1	1.3	6.8bTES	19	? 1				
	date	hrmn	sec	lat	long	depth	no	m	rms	damp	erln	erlt	erdp	
13	2	9	1416	8.59	059.46N	77 28.5W	149.5	28	3	1.16	0.000	21.6	17.2	5.7
	stn	dist	azm	ain	w	phas	calcphs	hrmn	tsec	t-obs	t-cal	res	wt	di
	LCO	3398	168.3	40.0	0	P	P	1422	9.6	360.99	359.70	1.29	1.00	10
	RCBR	4682	99.8	36.3	0	P	P	1423	45.0	456.39	457.12	-0.72	1.00	7
	PFO	5437	315.8	33.9	0	P	P	1424	39.4	510.86	510.52	0.35	1.00	7
	EFI	6121	165.4	31.7	0	P	P	1425	25.1	556.49	555.86	0.63	1.00	7
	EFI	6121	165.4		0	IVmB_BB		1425	28.3	559.7				
	EFI	6121	165.4		0	IAmb		1426	9.7	601.1				
	CMLA	6687	46.1	29.9	0	P	P	1425	59.1	590.53	591.50	-0.97	1.00	2
	CMLA	6687	46.1		0	IVmB_BB		1426	3.6	595.0				
	CMLA	6687	46.1		0	IAmb		1426	4.2	595.6				
	MACI	7140	59.0	28.5	0	P	P	1426	27.2	618.59	618.76	-0.17	1.00	2
	MACI	7140	59.0151	0.0	0	pP	pP	1427	3.7	655.14	654.18	0.96	1.00	5
	PAB	8504	49.6	24.3	0	P	P	1427	40.6	691.97	693.14	-1.17	1.00	1
	PAB	8504	49.6155	3.0	0	pP	pP	1428	18.2	729.56	729.75	-0.19	1.00	5
	PAB	8504	49.6	26.2	0	S	S	1437	15.41266	861266.68		0.18	1.00	8
	KDAK	9045	328.4	22.7	0	P	P	1428	6.6	717.97	719.55	-1.57	1.00	3
	BFO	9626	41.8	20.8	0	P	P	1428	32.9	744.36	746.02	-1.65	1.00	1
	BFO	9626	41.8		0	IVmB_BB		1428	37.2	748.6				
	BFO	9626	41.8		0	IAmb		1428	37.9	749.4				
	BFO	9626	41.8158	8.0	0	pP	pP	1429	11.1	782.48	783.45	-0.96	1.00	4
	KBS	9899	11.1	20.0	0	P	P	1428	45.8	757.23	757.65	-0.42	1.00	1
	KBS	9899	11.1		0	IAmb		1428	49.9	761.4				
	KBS	9899	11.1		0	IVmB_BB		1428	51.9	763.3				
	KBS	9899	11.1	35.6	0	PP	PP	1432	20.5	971.87	968.68	3.20	1.00	2
	KEV	10472	19.7	19.5	0	P	P	1429	8.0	779.44	781.44	-2.00	1.00	0
	KEV	10472	19.7		0	IAmb		1429	11.9	783.3				
	KEV	10472	19.7		0	IVmB_BB		1429	13.5	784.9				
	KEV	10472	19.7	26.5	0	SP	SP	1441	25.81517	181514.82		2.36	1.00	6
	BILL	10982	340.0	18.9	0	P	Pdif	1429	29.2	800.57	801.82	-1.25	1.00	2
	BILL	10982	340.0		0	IVmB_BB		1429	32.3	803.7				
	BILL	10982	340.0		0	IAmb		1429	35.1	806.5				
	BILL	10982	340.0	11.7	0	SKSac	SKSac	1439	53.61425	021425.40		-0.37	1.00	4
	KURK	13871	18.2	8.0	0	P	PKPdf	1434	49.31120	701121.34		-0.64	1.00	0
	NWAO	16181	202.1	13.2	0	P	PKPbc	1435	28.61159	981160.26		-0.28	1.00	3
	ENH	16499	348.5	6.8	0	P	PKPdf	1435	33.21164	591164.16		0.44	1.00	1
	KMI	17148	359.6	5.9	0	P	PKPdf	1435	42.81174	191173.26		0.94	1.00	1
	PALK	17448	69.0	5.4	0	P	PKPdf	1435	45.81177	261176.64		0.62	1.00	1
	CHTO	17812	10.0	4.8	0	PKPdf	PKPdf	1435	49.61181	041180.59		0.45	1.00	1
	CHTO	17812	10.0175	2.0	0	pPKPdf	pPKPdf	1436	28.51219	871220.40		-0.53	1.00	5
	CHTO	17812	10.0177	3.0	0	sPKPdf	sPKPdf	1436	45.41236	841235.95		0.89	1.00	10
	KAPI	18072	256.6	4.2	0	P	PKPdf	1435	52.21183	621183.05		0.57	1.00	2
	2013	2	9	1416	8.6	D	0.991	-77.475149.5	TES	19	1.2	6.8bTES	7.0BTES	

Output from the location program: Adding three pP phases and one pPKPdf and sPKPdf each, the hypocenter depth changes (shown in the third and last line) from 132 to 149 km. While this is now more different from the solution given by the USGS, we take this as the best solution determined from our data set, which may be even better than the USGS solution, because the latter commonly uses only P-wave first arrivals for their hypocenter location. Yet, the reading, proper identification and time picking of later arriving phases is crucial for significantly reducing the errors in hypocenter depth estimates (see Figure below).



Left: hypocenter locations using only P phases; **middle:** by adding S phases; **right:** by adding also depth phases and core reflections. (Copy of Figure 7 from IS 11.1, modified from Schöffel and Das, *J. Geophys. Res.*, Vol. 104, No. B6, page 13, Figure 2; © 1999, by permission of American Geophysical Union).

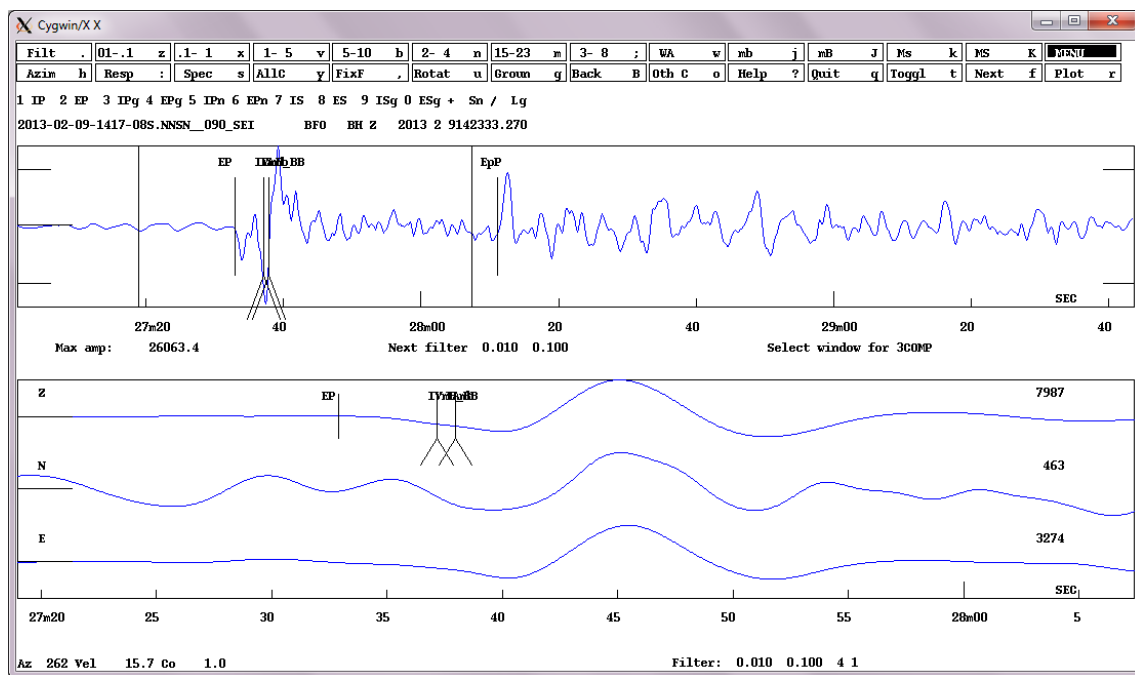


Map showing the epicenter of the Colombia earthquake (red star) determined by using a network of globally distributed stations (blue triangles). For comparison we plotted the epicenter from the single station (BFO, see label) 3-component location as a green star. For details of the BFO location see the next section.

6 Single station 3-component source location

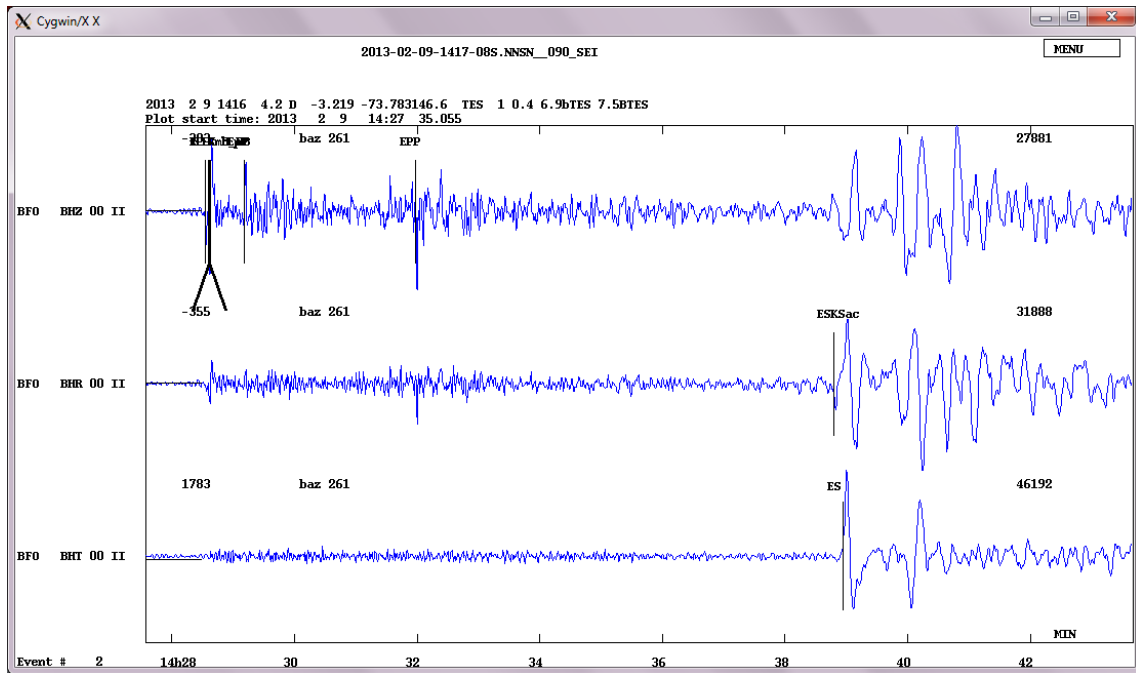
So far we have demonstrated how data from a global network can be processed in SEISAN. This is nowadays the common procedure as the data from a large number of stations are now freely accessible via Internet. However, it is also possible with SEISAN to estimate the event location by analyzing only the three-component records of a single station. The epicentral distance of the station can be determined best from the time difference S-P, but also using time differences of P to other later phases such as PP, SP or even SS (see Fig. 11.23 of Chapter 11 in Bormann2009, and EX 11.2). SKS-P is not suitable, because this time difference does not change much with distance. The backazimuth can be determined from the amplitude ratios of the three components (see Figure 1 and text in IS 11.1 and EX 11.2). For the backazimuth calculation the method of Roberts et al. (1989) is implemented in SEISAN.

We start with the determination of the back-azimuth from the P-wave first-motion amplitudes, as we need the backazimuth to rotate the seismograms into the ZRT system, which we will use to identify the phases.



Three-component analysis of station BFO at $\Delta = 86.4^\circ$, filtered 0.01-0.1 Hz: From the above trace plot we see a positive correlation between the vertical and the two horizontal components. By definition this means that first motion direction of P in the horizontal components shows toward the source, which is in the south-west. However, since the N-S amplitude is much smaller than the E-W amplitude (compare the numbers on the outermost right side of the traces), we can further conclude that the ray is coming more from the west than from the south thus matching for station BFO a backazimuth of 262° .

In the next step we estimate the epicentral distance via the time difference between P and identified later arriving phases.



Three-component analysis of station BFO at $\Delta = 86.4^\circ$, unfiltered: From the time difference between P with PP and S (here identified as SKS on the middle trace and as the slightly later arriving S on the bottom trace) we estimate the epicentral distance, assuming a source depth of 150 km according to the pP-P time (see Table below). This depth estimate agrees reasonably well with interpreting the strong onsets that follow SKS. S arrives about 13 s after SKS and the depth phase sS 74 s after S, corresponding at $\Delta = 86.4^\circ$ to a computed source depth of 170 km according to the IASP91 tables. This gives some feeling for the possible influence of errors in the assumed velocity model, phase interpretation and/or onset-time piking on the estimation of source depth.

We can either read the distance for which we get the best match between our observations and the travel time curves of the global IASP91 velocity model, or equivalently estimate it with a grid search by computing travel times for a range of distances as done here. In our example we then get $\Delta = 86.5^\circ$ (see Table below). This differs only 0.15° from our first multi-station P-wave location (86.35°) and even less from our improved multi-station and multi-phase location (86.56°) (see above location program outputs). Together with the backazimuth we can now locate the event on a suitable stereographic map with station BFO in the center (as in Figure 12 of EX 11.2 for an Ecuador earthquake and station CLL).

Phase difference	Time difference $\delta t(s)$	$\delta t(s)$ for $\Delta = 86^\circ$	$\delta t(s)$ for $\Delta = 87^\circ$
PP-P	205.7	202	205
S-P	623.3	622	626

```

#      2  9 Feb 2013 14:16  1  D                                ? 1

date hrmn  sec      lat      long depth  no m    rms  damp  erln  erlt  erdp
13 2 9 1416  4.23  312.63S 73 47.5W 146.6   5 2    0.37 0.000552.7488.0 0.0
stn  dist  azm  ain  w  phas  calcp  hrs  tsec  t-obs  t-cal  res  wt  di
BFO  9700 41.4 14.4 0  SKSac  SKSac  1438 47.31363.051362.11 0.94 0.98*12
BFO  9700 41.4 35.9 0  PP      PP      1431 58.6 954.40 954.80 -0.39 1.00*22
BFO  9700 41.4      0      1428 35.3 751.1
BFO  9700 41.4159.1 0  pP      pP      1429 11.1 786.84 786.42 0.42 1.00*22
BFO  9700 41.4      0  IAmb      1428 37.9 753.7
BFO  9700 41.4      0  IVmB_BB  1428 37.2 753.0
BFO  9700 41.4 20.4 0  P      P      1428 32.9 748.67 749.64 -0.97 0.97*11
BFO  9700 41.4      AZ      262.0 262.0 0.00 0.20 33

2013  2 9 1416  4.2 D  -3.210 -73.792146.6  TES  1 0.4

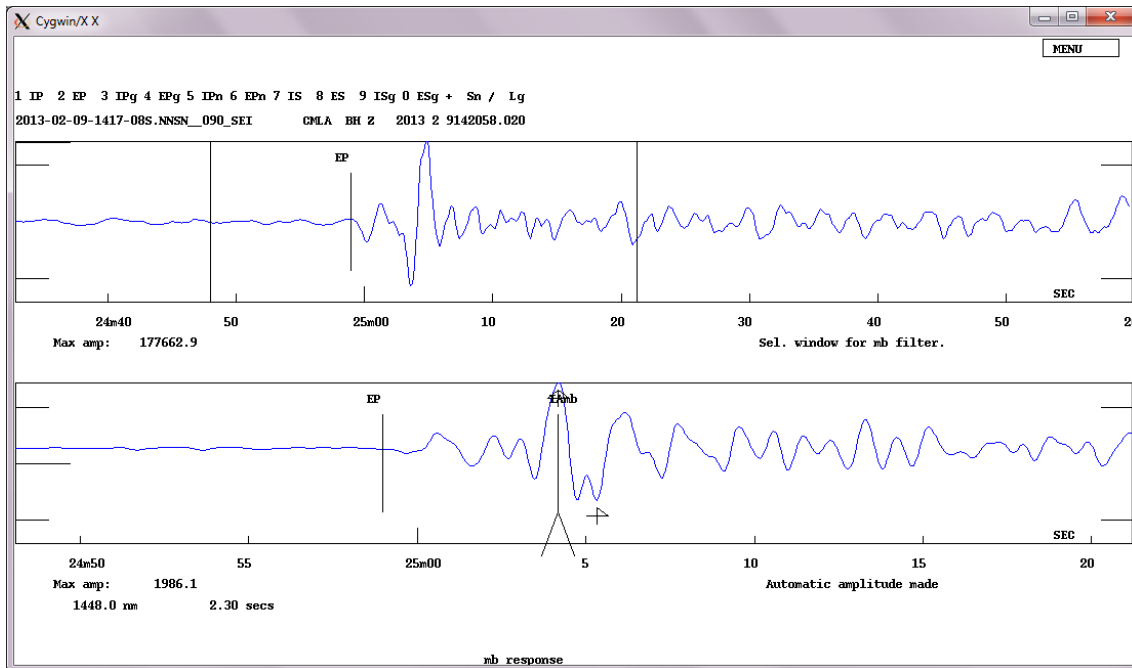
```

Output from the location program: We can obtain a single-station three-component location, which is within 500 km from the one obtained by the standard location program with several stations. While our distance estimate is very good, we would have to get a 5° larger back-azimuth reading to get a location very close to the multi-station location. This shows the approximate range of uncertainty of single station locations in the far teleseismic range. Wrong phase interpretation may increase single station location errors. E.g., when interpreting in this case the earlier arriving SKS as S, the distance estimate from S-P would be 0.2° shorter in the case of BFO but this underestimation would increase for any station further away within the distance range where SKS arrives before S (see differential travel-time curve in Figure 4 of EX 11.2). For shallow earthquakes at $\Delta = 100^\circ$ the error $\delta\Delta$ would reach already 0.9°. However, with proper S-P times, or when several other phases have also been correctly identified, the distance can generally be estimated, up to 160°, within 1-2° and the azimuth with P-wave first motion amplitudes, when measured with good signal-to-noise ratio on well calibrated 3-component seismometer components, within about 5-10°. This allows rather fast and reasonably good first location estimates even with single station data only (see Fig. 11.23 in Chapter 11 of NMSOP-1 in Bormann, 2009, and examples in EX 11.2).

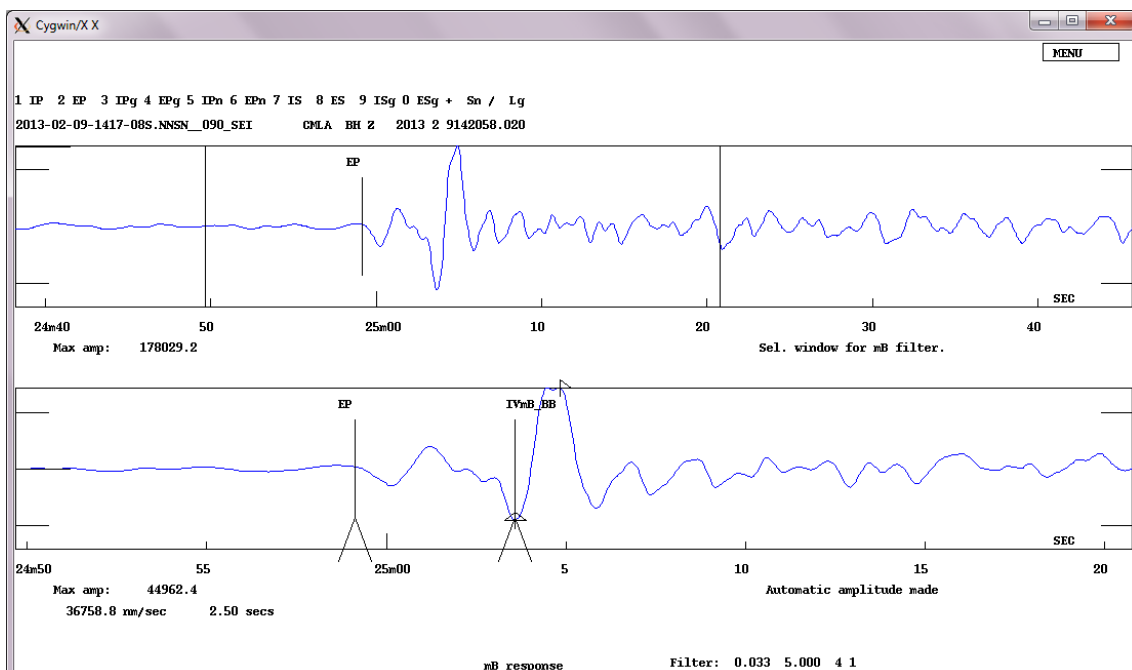
Finally, we determine the teleseismic magnitudes for this earthquake. Since this is a deep earthquake, only the two teleseismic standard body-wave magnitudes m_b and m_{B_BB} should be calculated. Note, however, that the IASPEI WG on magnitudes encourages stations to measure according to the standard procedures also IAMs_20 and IVMs_BB with their periods if surface-waves are still clearly visible in records of deeper earthquakes as this is the case also for our example (see, e.g., in section 5, the record trace for station KADAK in the plot: *Vertical-component traces for which phases are picked, filtered 0.02-5 Hz*). Such data are of great interest for investigating the depth and period dependence of surface waves and to derive improved calibration functions $\sigma(\Delta, h, T)$. Note that for 20 s surface-waves and thus M_{s_20} calculation such depth corrections have already been published by Herak et al. (2001):

$$\begin{aligned}
 \Delta M_s(h) &= 0 && \text{for } h < 20 \text{ km} \\
 \Delta M_s(h) &= 0.314 \log(h) - 0.409 && \text{for } 20 \text{ km} \leq h < 60 \text{ km} \\
 \Delta M_s(h) &= 1.351 \log(h) - 2.253 && \text{for } 60 \text{ km} \leq h < 100 \text{ km} \\
 \Delta M_s(h) &= 0.400 \log(h) - 0.350 && \text{for } 100 \text{ km} \leq h < 600 \text{ km}.
 \end{aligned}$$

For station KDAK, e.g., one would measure $IAMs_{20} = 10078$ nm at $T = 19$ s and thus get a depth corrected $M_{s_20} = 6.2 + 0.5 = 6.7$ which compares rather well with our body-wave magnitudes $m_b = 6.8$ and $m_{B_BB} = 7.0$ determined in the following.



Amplitude reading for mb magnitude at station CMLA at $\Delta = 59.9^\circ$: As in the case of the regional event, the upper broadband trace is converted into a WWSSN short period trace (lower trace) on which the amplitude Iamb is read, here as 1148 nm with a period of 2.3 s.



Amplitude reading for mB_BB magnitude at station CMLA at $\Delta = 59.9^\circ$: As in the case of the regional event the amplitude IVmB_BB is read on the bottom trace that has been deconvolved into a broadband velocity trace (top trace shows uncorrected data, which are in the case of a broadband seismometer also proportional to velocity): here 36758 nm/s with a period of 2.5 s.

Computing the average magnitudes from all station amplitude observations we get the values below. In this case mB_BB is equal to Mw, and mb is only slightly smaller. Since the period at which mb is read is almost the same as for mB_BB the difference is mainly due to the larger bandwidth of the broadband record.

Magnitude scale	Value
Mw (USGS)	7.0
mb	6.8
mB_BB	7.0

The following station amplitude readings were used in the event magnitude computation:

EFI	BZ	dist:	6121.0	amp:	13012.5	T:	2.7	mB =	6.9
EFI	BZ	dist:	6121.0	amp:	788.3	T:	1.1	mb =	6.5
CMLA	BZ	dist:	6687.0	amp:	36756.6	T:	2.5	mb =	7.9
CMLA	BZ	dist:	6687.0	amp:	36758.8	T:	2.5	mB =	7.5
CMLA	BZ	dist:	6687.0	amp:	16414.7	T:	2.3	mb =	7.6
KEV	BZ	dist:	10472.0	amp:	83.5	T:	0.8	mb =	6.0
KEV	BZ	dist:	10472.0	amp:	2421.2	T:	5.5	mB =	6.6
KBS	BZ	dist:	9899.0	amp:	500.9	T:	0.9	mb =	6.5
KBS	BZ	dist:	9899.0	amp:	11474.4	T:	7.2	mB =	7.1
BFO	BZ	dist:	9626.0	amp:	34897.1	T:	3.7	mB =	7.4
BFO	BZ	dist:	9626.0	amp:	1747.1	T:	1.1	mb =	6.9
BILL	BZ	dist:	10982.0	amp:	1555.7	T:	6.0	mB =	6.7
BILL	BZ	dist:	10982.0	amp:	106.9	T:	1.2	mb =	6.3

7 Concluding remarks

The examples given here are meant to illustrate how the data processing is done using the SEISAN software and to illustrate some basic concepts of earthquake data processing. The examples should help newcomers to earthquake data processing to get started. It will be necessary to first learn SEISAN, but then the data used can be downloaded and the steps shown here followed. When analyzing seismograms, the number of phases that can be read depends on the earthquake size and thus the SNR of the recording, the epicentral distance, the source depth as well as on the bandwidth and period range of the seismic records (see Figs. 2.57-2.61 in Chapter 2). From a scientific perspective we wish to have as many good phases read as possible. However, there is often a limitation of how much can be done due to time constraints. We encourage the reader to read more than the first arriving P, and at least to identify clear additional later phases. A good exercise is to download data from the IRIS Wilber, take the data into SEISAN and start processing.

Acknowledgments

Our thanks go to Jens Havskov, Won-Young Kim and James Dewey. Their review comments helped to improve this manuscript.

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